

Integrated Aeromagnetic and Aero-radiometric Data for delineating Lithologies, Structures and Hydrothermal Alteration Zones in Ganmo–Amoyo–Ojoku Areas, Nigeria

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ABSTRACT

Aeromagnetic and aero-radiometric methods have been successfully used to locate zones of mineralization potentials in Ganmo–Amoyo–Ojoku areas by mapping lithology, structures and hydrothermal alteration regions for industrial purposes. The data, sheet 223-Ilorin on a scale of 1:100,000 were obtained from the Nigerian Geological Survey Agency (NGSA) at a height of 80 meters along a series of NE-SW flight lines spaced 500 m interval. The Total Magnetic Intensity data was gridded at 100 m then reduction to the magnetic equator (RTE) was applied to center magnetic anomalies over their respective magnetic source bodies with the aid of Oasis Montaj™ software. Other enhancement techniques applied are first vertical derivative (FVD), analytic signal (AS) and tilt angular derivative (TDR). In addition, the potassium to equivalent thorium (K/eTh) ratio and ternary radiometric maps were revealed to support the delineation process. The interpreted map for magnetic lineaments by delineating the anomaly maxima mostly observed in the southern and eastern portions of the study area extending from Orogun to Apala and Alaya with continuous linear contour patterns. The lineaments range in lengths from 50 m to 1000 m while the orientations range from NE-SW to NW-SE. The equivalent ratio of potassium-thorium and ternary map revealed lithological units and hydrothermal alterations within the basement complex. From this study, volcanic intrusions, lineament structures, and hydrothermal alteration zones were delineated. These intrusions resulted from the high content of magnetic anomalies related to the mineralized potential zones in the area. The integration of both magnetic and radiometric maps has revealed an approximate understanding between radiation zones, observed fractures, and faults in the locations between the boundaries of rock types in the study area.

Received: 19th July 2026

Accepted: 23rd July 2026

Published: 31st July 2026

Keywords:

Aeromagnetic data, radiometric data, structural lineaments, Oasis Montaj, hydrothermal alteration

INTRODUCTION

Many authors emphasized the importance of lineament mapping as part of various geological investigations (e.g. Mwaniki *et al.*, 2015; Ramli *et al.*, 2010; Magaji & Sanusi, 2021). Lineament extraction can help interpret the regional structure of a given area and discern its different structural elements (Magaji and Sanusi, 2021; Abdelouhed *et al.*, 2022; Bentahar *et al.*, 2020; Fossi *et al.*, 2021). Lineaments can also assist in understanding underlying rocks architecture (e.g. basement faults), exploring mineral deposits (Magaji *et al.*, 2025a; Adepoju *et al.*, 2021; Hubbard *et al.*, 2012; Liu *et al.*, 2000), and discovering groundwater reservoirs (Al-Saedi, 2013; Bruning *et al.*, 2011; Saint Jean Patrick Coulibaly *et al.*, 2020; Adhab, 2014; Soliman & Han, 2019). The description given by Benabbas (2006) was retained, defining the lineaments as rectilinear or slightly curved structures that appear in landscape and which may correspond either to physiographic features linked to vegetation, hydrography, or morphology: ridges, relief limits, outcrop lines, or to elements of structural origin: faults, reversals of layers, a dense network of joints.

In digital image processing, lineaments usually occur as edges with tonal differences. Most extraction techniques are based

on edge enhancement by improving the tone, the hue, and the image structure (Magaji *et al.*, 2025b; Folorunso *et al.*, 2024; Hung *et al.*, 2005; Javhar *et al.*, 2019). In this regard, two methods are proposed for identifying and extracting lineaments from remote sensing data: manual and automatic. The manual method consists of enhancing the tonal differences, then visually interpreting the lineament segments of the satellite image. This technique is efficient but it is relatively time-consuming and more importantly highly subjective. The automatic method uses computer software and algorithms to detect and extract lineaments automatically. Although this technique is quick, it generates many insignificant lineaments and must go through a post-processing phase that includes validation and removing anthropic lineaments. As a result, some manual processing should be incorporated (Abdelouhed *et al.*, 2022). However, Ahmadi and Pekkan (2021) proposed that a single method may not provide an accurate lineament distribution. It is therefore necessary that the identification and extraction of lineaments integrate different methods.

This study thereby combines magnetic and radiometric methods to detect and extract lineaments which reflect deep geological structures such as basement faults, folds, joints etc.

Magnetic method is used in the investigation of the subsurface geology of an area. This is achieved by detecting magnetic anomalies within the earth's magnetic field which are normally caused by magnetic susceptibility changes of the underlying rocks. Most rock forming minerals are non-magnetic but a few that contain sufficient amount of magnetic grains in them can impact magnetism to their host rock thus, producing detectable magnetic anomalies. The application of radiometric method is useful in geologic mapping because of its ability to detect radioactivity emanating from the natural decay of elements such as uranium, thorium and potassium from the earth material. The product of this decay results in the emission of alpha, beta and gamma radiations which can be detected on the surface using radioactive element detectors.

Magnetic and radiometric data have been conjunctively used in the study of lithologies and structures in Yola area, Nigeria (Bassey and Ishaku, 2012), in Ashanti gold belt of Ghana (Boadi *et al.*, 2013). Radiometric exploration for uranium in N.E. Nigeria has been done by several researchers (Okujeni, 1987; Okujeni *et al.*, 1987; Mustapha 1992; Su and Dada, 1998; Haruna *et al.*, 2011). The present work jointly analyze data from aeromagnetic and ground radiometric sources over

Ganmo-Amoyo-Ojoku areas. Hence information from both aeromagnetic and ground radiometric data can be well correlated with ground information. The objectives of this work are to identify structural features, lithologies, magnetic source and hydrothermal alteration zones from images of the area.

Geographic Location and Geological Settings of the Study Area

The study area (Figure 1) is located within Ilorin Sheet 223 NW in southwestern Nigeria and forms part of the Nigerian Basement Complex, a segment of the Pan-African mobile belt. This basement terrain is composed predominantly of high-grade metamorphic rocks and granitoids that have undergone multiple phases of deformation, metamorphism, and magmatism during the Pan-African orogeny. Regionally, the Nigerian Basement Complex is subdivided into the migmatite-gneiss complex, the schist belts and the Pan-African granitoids. The migmatite-gneiss complex constitutes the oldest component and comprises banded gneisses, migmatites, and associated granitoid intrusions (Ajibade *et al.*, 1987; Dada, 1998; Turner, 1983).

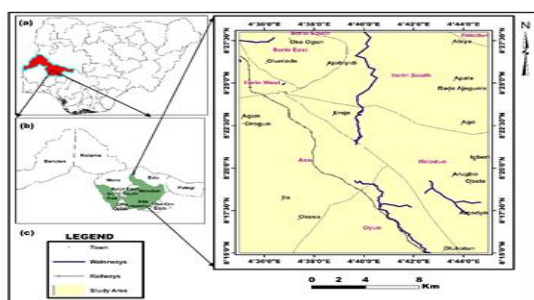


Figure 1: Location map of the study area

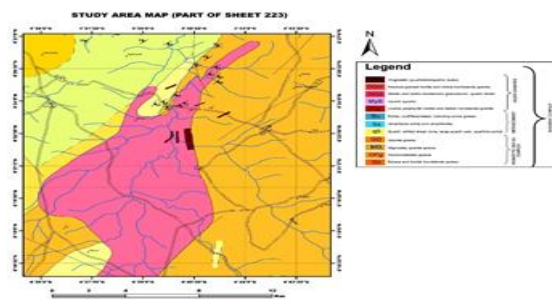


Figure 2: Geological map of the study area

The Pan-African granitoids, often referred to as Older Granites, are widely distributed and range in composition from dioritic to granitic (Rahaman, 1989). These rocks are commonly associated with late-stage pegmatitic and aplite intrusions, reflecting the final stages of magmatic evolution. Within the Ganmo-Amoyo-Ojoku geological map (Figure 2), the principal lithologies include biotite gneiss, augen gneiss, charnockitic rocks, porphyritic diorites, and granitoids, with pegmatite and aplite occurring as cross-cutting bodies. Field relationships suggest that the gneissic rocks form the basement framework into which younger intrusive bodies were emplaced. Pegmatitic veins commonly cut across earlier lithologies, indicating emplacement during the late stages of magmatic activity. Overall, the geological setting reflects a structurally complex basement terrain in which lithological diversity and deformation patterns are closely linked (Rahaman, 1988). This context is essential for understanding the geophysical signatures observed in the study area and for developing an integrated interpretation of subsurface architecture.

MATERIALS AND METHODS

Data Acquisition

The aeromagnetic and aero-radiometric data used for this work is sheet number 223 bounded between longitudes 004° 00' E to 004° 30' E and latitudes 08° 30' N to 09° 00' N, on a scale of 1:100,000. The data were acquired as part of a second nationwide aeromagnetic survey conducted in two phases by Fugro Airborne Surveys from 2005 to 2010 on behalf of the Federal Government of Nigeria (FGN). The data were

acquired (i) along a series of NW – SE flight lines at an interval of 500 m, and a tie line spacing of 2000 m along NE-SW direction, (ii) at an average of 80 m terrain clearance, and (iii) data recording interval of 0.1 s (~7 m). The new data has an added advantage of high resolution over the data acquired in the first nationwide survey conducted between 1974 – 1976, which were obtained using a flight elevation of 150 m and flight line spacing of 2000 m.

Data Processing

Reduction to the equator (RTE) correction was performed on the aeromagnetic anomaly data (Figure 3) because of the variation in inclination and declination of the geomagnetic field from the geomagnetic equator to the pole, such that the peaks of total magnetic anomalies due to subsurface are shifted from directly over the source and the anomalies becomes highly asymmetric. By applying this correction, the spatial shift in the shape and the peak of the magnetic anomalies over their causative sources are corrected (Hinze *et al.*, 2013). Thus, a RTE correction of 4.132° declination, 9.5342° inclination, a grid cell size of 100 m, and a minimum curvature gridding was applied for this area using the Oasis Montaj (R) software, version 8 (Phillips, 2007; Lawal and Nwankwo, 2017). For the radiometric data, potassium (K), thorium (Th), and uranium (U) were procured in grid form and processed using the Oasis Montaj (R) software, version 8.

First Vertical Derivative (FVD)

The FVD is used to enhance shallow features with their boundaries and associated lineaments (Eldosouky *et al.*, 2017). This technique has the ability of amplifying high-wave number anomalous features at the expense of the lower ones (Ibraheem *et al.*, 2019). As a result, shallow features can be properly identified and mapped (Nabighian 1972). Evjen

(1936) was the first to delineate edges of gravity and magnetic bodies using the first vertical derivative. If M is the magnetic field, the expression for the FVD is given as (Ibraheem *et al.*, 2019).

$$\text{FVD} = \frac{\partial M}{\partial x} \quad (1)$$

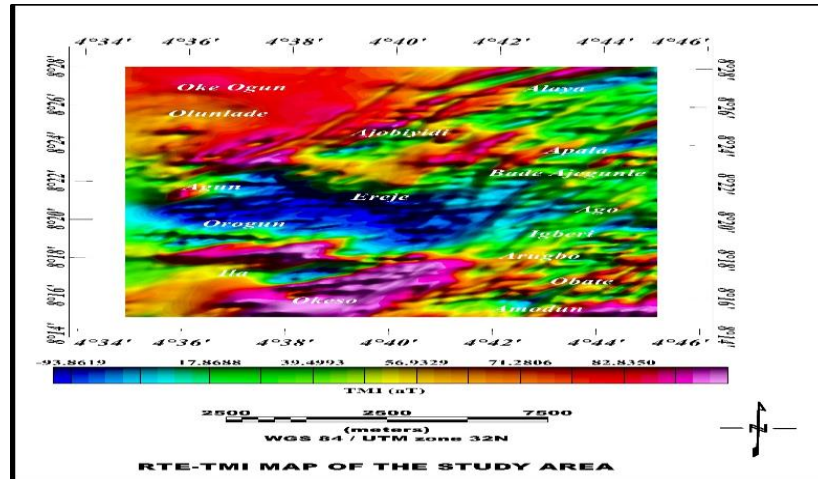


Figure 3: RTE-TMI map of the study area

Analytic signal (AS)

Analytic signal is an enhanced technique which is related to potential fields (e.g., magnetic field data) by obtaining their derivatives. The expression is given as follows (Nabighian 1972):

$$\text{AS}(x,y,z) = \sqrt{(\partial M/\partial x)^2_x + (\partial M/\partial y)^2_y + (\partial M/\partial z)^2_z} \quad (2)$$

Where $\hat{x}$, $\hat{y}$ and $\hat{z}$ are the unit vectors in x , y , and z directions

Tilt Angle Derivative (TDR)

The tilt derivative filter has been demonstrated by Verdusco *et al.*, (2004) to carry out an automatic gain-control (AGC) function, which attempts to balance the response from both strong and weak anomalies. As a result, the filter offers a useful method for locating notable anomalies within the study region. The TDR was performed on the Residual Magnetic Anomaly (RMA) maps. According to Miller and Singh (1994), the tilt angle is defined as:

$$\theta = \text{TDR} = \tan^{-1} \frac{\text{FVD}(z)}{\text{HGM}(x,y)} \quad (3)$$

Where $\text{FVD}(z)$ = First Vertical Derivative and $\text{HGM}(x,y)$ = Total Horizontal Gradient

Although the tilt angle gradient is a dimensionless ratio, it responds well to both deep and shallow sources, as well as sources at the same level with a wide dynamic range of amplitudes (Cooper, 2006). The TDR values are much easier to grasp because they are limited to values between $-\delta/2$ and $+\delta/2$ (Cooper, 2006).

RESULTS AND DISCUSSION

First Vertical Derivative map

The results obtained after applying the FVD filtering method (Figure 4) give a good representation of features closer to the subsurface, e.g., faults (white lines). In this case, the filtration

of data enhanced high-frequency anomalies so that shallower sources could be mapped at the expense of the low-frequency ones. Almost all the entire area is occupied by these high-frequency anomalies, indicating shallow depth to causative sources in comparison with some other parts such as southwest, northeast, northwest and part of north central with deeper sources which have the presence of sedimentary covers (Salem *et al.*, 2007). Also, on the map are long and short linear features distinguishing the edges of magnetic source bodies with zero contour lines representing the locations of the contacts or faults. The long linear features, which are continuous and less prominently trend in the NE-SW and NW-SE directions, represent lithological boundaries and could be interpreted as major fracture/fault lines on the map, while the short linear and curvilinear features with their orientation trending mostly NW-SE and E-W are also concentrated within the lithological boundaries of magnetic source bodies of migmatite gneiss and biotite gneiss. It is also interesting to note that both long and short linear features are defined by a contour value of zero that coincides with the peaks in both AS map (Figure 5).

Analytical Signal (AS) map

The AS map (Figure 5) reveals magnetic anomalies ranging from 0.0065 to 0.2249 nT/km. The map was interpreted for magnetic lineaments by delineating the anomaly maxima mostly observed in the southern and eastern portions of the study area extending from Orogun to Apala and Alaya with continuous linear contour patterns. The lineaments range in lengths from 50 m to 1000 m while the orientations range from NE-SW to NW-SE. In the first magnetic domain, the majority of maxima are trending in the NE-SW direction, while minor maxima are observed to be trending in the NW-SE direction.

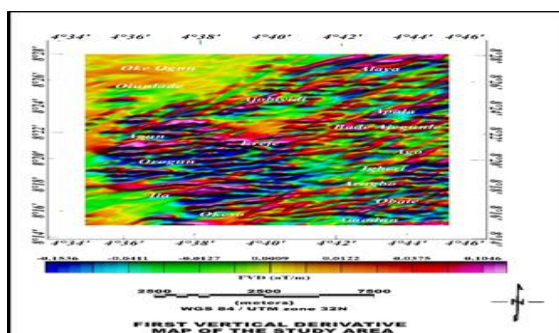


Figure 4: FVD map of the study area

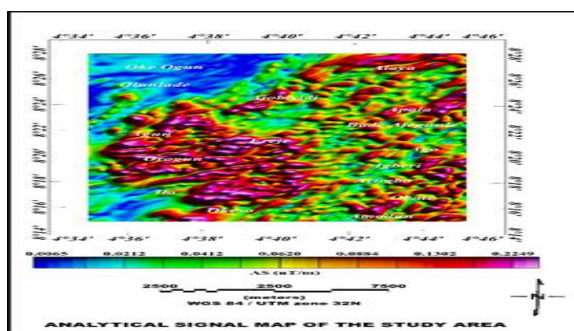


Figure 5: AS map of the study area

These two groups of maxima can be associated with intruded granite bodies (e.g. biotite granite, granite gneiss) rich in ferromagnesian-bearing rocks with a low amount of felsic minerals (Telford *et al.*, 1990; Balogun, 2019; Folorunso *et al.*, 2024; Magaji *et al.*, 2025). In contrast, the low- and intermediate-magnetic intensity zones might be reflections of downfaulted blocks of the basement complex filled with sediments or possibly a set of tectonic trends controlling the overlying sediments which can be observed at Oke Ogun and Olunlade areas.

The Tilt Derivative (TDR)

The TDR filter is normally utilized in the detection and mapping of geologic structures which can be related to shallow basement structural features or mineral exploration targets. The TDR filter was applied to the RTE anomaly data, to provide greater resolution of structural features. The TDR map (Figure 6) demarcates possible edges of geologic source bodies. The TDR technique facilitates the horizontal location with the extent of edges of magnetic source bodies within the study area. When utilizing the RTE transformation, the TDR is negative (blue and green color) when it is over magnetic

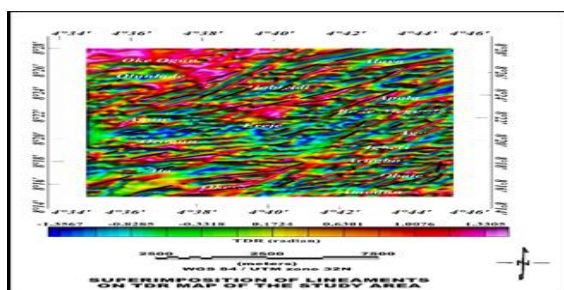


Figure 6: TDR map of the study area

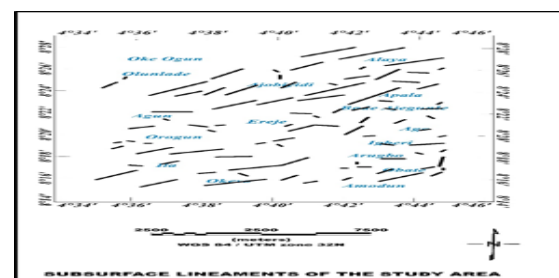


Figure 7: Subsurface Lineaments map

The K/eTh map

As high % K and low concentration of eTh are associated with alterations in many ore deposits (Ostrovskiy 1975), the K/eTh ratio map is of great importance for searching signatures associated with hydrothermal alteration zones. This map is depicted in Figure 8. Therein, the areas affected by hydrothermal process are revealed in red to pink color (i.e. 0.2302 to 0.5956 %/ppm) (strong indicator, SI). These SIs are found in the southern, central and northern regions of the study area i.e. Illa, Orogun, Ereje, Ajobiyyidi and Oke Ogun though more prominent in the southwest and northeast. These areas are associated with amphibolite, migmatite gneiss, biotite gneiss and porphyritic diorites. On the other hand, the low indicator (LI) observed on the map shows that potassium is low and the thorium concentration is high. As observed on the map, most of these alteration zones are trending towards the NE-SW direction and follow the same pattern with the aeromagnetic data enhancement techniques.

The ternary map

Finally, the ternary map (Figure 9) of the study area was generated by modulating the three colours red (R), green (G), and blue (B) in equal proportion so that their intensities could be revealed. From this map, the lithological contrasts resulting from geological and mineral explorations can be interpreted. Here the white colour found in some areas across the study area shows the high level of concentration (HLC) for the radioactive elements which is characterized by strong radiometric response, and these areas can be associated with amphibolite, migmatite gneiss, biotite gneiss. The low concentration of the radioactive elements that represented black (LC) is a resultant of porphyritic diorites. The red colour represents high K content but low Th and U related to granitic body, while the green colour corresponds to regions of high Th, low K and U associated with metamorphic rock formations in the study area.

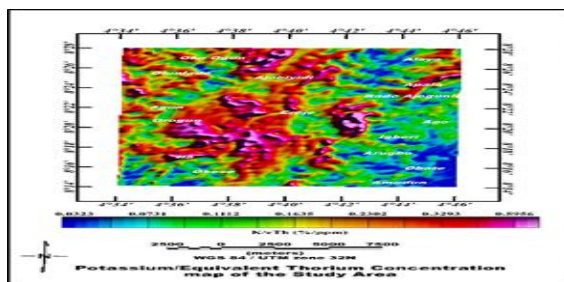


Figure 8: K/eTh map of the study area

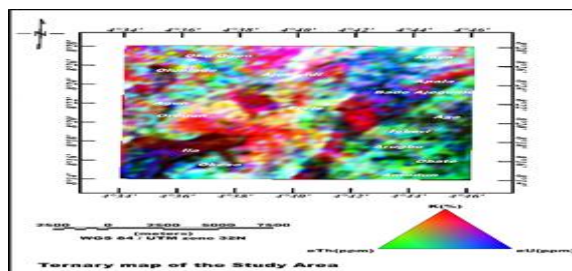


Figure 9: Ternary map of the study area

CONCLUSION

Aeromagnetic and aero radiometric maps were used to evaluate the spatial relationship between the subsurface structural features, alteration zones, and various lithologies in the zone located in the Ganmo–Amoyo–Ojoku areas of the Migmatite–Gneiss complex units of the southwestern basement complex of Nigeria. The FVD map showed a good representation of features closer to the subsurface by enhancing high-frequency anomalies so that shallower sources could be mapped at the expense of the low-frequency ones. The analytic signal map also showed networks of magnetic discontinuities that might represent fracture zones that could be targets for mineral prospection. The structural features obtained from the magnetic data using various enhancement techniques show that NE, NW, ENE, and NS are the main directions in the area. Evaluation of radiometric data further revealed various lithologies (migmatite gneiss, biotite gneiss, amphibolite, pegmatite and granite diorite) and hydrothermal alteration zones trending in the NE, NW, ENE, and NS directions. Integration of both aeromagnetic and aero radiometric data has further helped in detecting areas where mineral resources can be explored.

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