



THE GROWTH FACTOR AND BULK HYGROSCOPICITY OF ATMOSPHERIC SOOT OF URBAN AEROSOLS

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ABSTRACT

Aerosols within urban atmosphere can be composed of water-soluble aerosols from industrial emissions, insoluble and soot from biomass and bio-fuel emissions respectively. In this study, simulation was carried out using Optical Properties of Aerosols and Clouds (OPAC) to model the hygroscopic growth factor and bulk hygroscopicity of Soot at spectral range of 0.25 to 1.00 μm for eight different relative humidities. The results in this study revealed that the aerosol hygroscopic growth factor increases with relative humidity (RH) while the bulk hygroscopicity decreases with increase in RH from 50-99% RHs. The aerosol hygroscopic growth factor increases with increase in RH while the bulk hygroscopicity decreases with increase in RH for the number, volume and mass ratios. The aerosol growth factor revealed that the mixture is barely hygroscopic, less hygroscopic and more hygroscopic from 50 – 80% RH, 90 – 95% RH and 98 – 99% RH respectively for the number mix ratio. The aerosol growth factor revealed that the mixture is less hygroscopic, more hygroscopic and most hygroscopic from 50 – 80% RH, 90 – 95% RH and 98 – 99% RH respectively for the volume and mass mix ratios. The bulk hygroscopicity ranges between 0.02007 to 0.09456 for the number mix ratio from model 1 to model 3, the bulk hygroscopicity ranges between 0.13596 to 0.32956 for the volume mix ratio from model 1 to model 3 while the bulk hygroscopicity ranges between 0.12831 to 0.29925 for the mass mix ratio from model 1 to model 3.

Keywords: bulk hygroscopicity, hygroscopic growth factor, OPAC, relative humidity, soot

INTRODUCTION

The microscopic solid or liquid particles suspended in the atmosphere are called aerosols, these aerosols have effect on the radiative balance of the Earth and thus, the climatic system by interacting directly with solar and terrestrial radiation or changing the formation of clouds indirectly (Lohmann and Feichter, 2005; IPCC, 2013 ; Seinfeld *et al.*, 2016). These aerosol climatic effects are highly irregular due to of the large variability of the physical and optical properties of aerosol, which are attributed to multiplicity of sources, and their dependence on the prevailing meteorological and atmospheric conditions (Satheesh and Krishna Moorthy, 2005). The aerosol optical properties are strongly dependent on relative humidity (Tijjani and Akpootu, 2013a). The strong effect of aerosols on climate has not been given significant attention, which present significant uncertainty into climate predictions (IPCC, 2007).

The Smog and Soot which are also referred to as ground-level ozone and particulate matter respectively are the two most common forms of air pollution (Meseke *et al.*, 2022). Soot aerosols also known as black carbon (BC) are formed when incomplete combustion takes place. The major sources of soot aerosols are forest fires, diesel engines and biomass burning. Absorption properties of soot particles depend highly on the combustion temperature and other material (e.g. organic carbon) emitted during the processes of combustion (Bond and Bergstrom, 2006). Soot is made up of monodispersed spherical particles that collect into mass fractal aggregates having a broad size distribution, the primary soot particles are usually very small (Tijjani and Akpootu, 2013b). The role of soot particles in combustion is the major rationale of both experimental and theoretical investigation of soot radiative properties (Akpootu and Momoh, 2013a). The dust aerosol

present in the atmosphere during the harmattan season in the northern hemisphere is a common feature of the climate of most parts of West Africa (Akande *et al.*, 2013). The enormous amount of dust and sand particles raised and transported by the harmattan dust haze strongly decreases visibility and are estimated to reach about 6000 m above sea level (Essienimo *et al.*, 2016a; Essienimo *et al.*, 2016b)

The hygroscopic growth and the mixing state of aerosol particles play a significant role for various atmospheric effects like the direct aerosol effect on climate, visibility degradation, and cloud formation (Sloane and Wolff, 1985; Pandis *et al.*, 1995; McFiggans *et al.*, 2006). Through the interaction of atmospheric particles and cloud droplets with incoming shortwave radiation, the particle hygroscopic growth is one of the major parameters influencing the terrestrial radiation budget and climate (IPCC, 2007). Some uncertainties connect to the hygroscopic growth estimation and cloud condensation nuclei (CCN) activation for the complex chemical mixtures of aerosol particles found in the atmosphere. Other uncertainties relates to how hygroscopic growth and CCN activation can be parameterized for implementation in higher scale climate models. The scattering and reflection of solar radiation by aerosols and clouds tends to cool the earth's surface, and this referred to as negative forcing while in a situation when the absorption of terrestrial radiation by greenhouse gases and clouds tends to warm it is referred to as positive forcing (Akpootu and Momoh, 2013b; Akpootu and Sharafa, 2013; Essienimo *et al.*, 2015a). The size distribution of any particular suspended particle determines the life-span of the particle in the atmosphere and the distance it can travel (Essienimo *et al.*, 2015a; Essienimo *et al.*, 2015b).

Numerous studies have been carried out to investigate the effect of atmospheric aerosols. In the paper of Tijjani and Akpootu (2012), they modeled the optical depths, asymmetry parameters and single scattering albedos of urban aerosols using Optical Properties of Aerosols and Cloud (OPAC) at spectral range of 0.25 μm to 1.0 μm for eight different relative humidities (RHs). The radiative forcings (RF) and Ångström parameters was computed from the obtained data. Based on the RF, they found that as the RH increases there is a small increase in warming from 0 to 70% but as from 80 to 99% RH there is an increase in cooling from the first to the third model. Akpootu and Gana (2013) modeled the hygroscopicity properties of water soluble aerosols component based on microphysical properties of urban aerosols using OPAC to determine the effect of relative humidity on hygroscopic growth factor and bulk hygroscopicity at spectral range of 0.25-1.00 μm . Akpootu and Abdul salami (2013) describes the hygroscopicity properties of water soluble aerosols component based on optical and microphysical properties of urban aerosols using simulated data obtained from OPAC to

determine the density mix ratio resulting from hygroscopic growth factor and bulk hygroscopicity at spectral range of 0.25-1.00 μm for eight different relative humidities (RHs). They found that the density mix ratio indicates that there is a steady increase in aerosol hygroscopic growth factor with RHs and decrease in the magnitude of bulk hygroscopicity. Other studies include Akpootu and Muhammad (2013), Akpootu and Tijjani (2014), Seinfeld *et al.* (2016) to mention but a few.

The aim of this paper is to investigate the effect of hygroscopic growth factor and bulk hygroscopicity of soot in relation to eight different relative humidities (0, 50, 70, 80, 90, 95, 98 and 99%) of urban aerosols using extracted microphysical properties of number mix ratio, volume mix ratio and mass mix ratio simulated from Optical Properties of Aerosols and Clouds (OPAC).

MATERIALS AND METHODS

The models extracted from OPAC are given in table 1.

Table 1: Compositions of aerosols types (Hess *et al.*, 1998).

| Components | Model 1 | Model 2 | Model 3 |
|---------------|---------------------------------|---------------------------------|---------------------------------|
| | No.density (cm^{-3}) | No.density (cm^{-3}) | No.density (cm^{-3}) |
| Insoluble | 1.5 | 1.5 | 1.5 |
| Water soluble | 20,000.00 | 20,000.00 | 20,000.00 |
| Soot | 110,000.00 | 120,000.00 | 130,000.00 |
| Total | 130,001.50 | 140,001.50 | 150,001.50 |

The urban aerosols data used in this study are derived from the Optical Properties of Aerosols and Clouds (OPAC) data set (Hess *et al.*, 1998). A mixture of three components was used to describe Urban aerosols: a water soluble (WASO) components consist of scattering aerosols that are hygroscopic in nature, such as sulphates and nitrates present in anthropogenic pollution, water insoluble (INSO) and Soot. The particle number densities of soot were varied as 110,000 120,000 and 130,000 cm^{-3} while the water soluble and insoluble components were kept constant.

The key parameter used to characterize the hygroscopicity of the aerosol particles is the aerosol hygroscopic growth factor $gf(\text{RH})$, which indicates the relative increase in mobility diameter of particles due to water absorption at a certain RH and has been defined as the ratio of the particle diameter at any RH to the particle diameter at $\text{RH} = 0$, the RH is taken for seven values 50%, 70%, 80%, 90%, 95%, 98% and 99% (Swietlicki *et al.*, 2008; Randles, *et al.*, 2004; Akpootu and Gana, 2013):

$$gf(\text{RH}) = \frac{D(\text{RH})}{D(\text{RH}=0)} \quad (1)$$

The $gf(\text{RH})$ are subdivided into different classes with respect to hygroscopicity. One classification is based on diameter growth factor by Liu *et al.* (2011) and Swietlicki *et al.* (2008) as barely Hygroscopic ($gf(\text{RH}) = 1.0 - 1.11$), less Hygroscopic ($gf(\text{RH}) = 1.11-1.33$), more Hygroscopic ($gf(\text{RH}) = 1.33-1.85$) and most hygroscopic growth ($gf(\text{RH}) > 1.85$).

Most of the atmospheric aerosols are externally mixed with respect to hygroscopicity, and consist of more and less hygroscopic sub-fractions (Swietlicki *et al.*, 2008). The ratio between these fractions as well as their content of soluble material determines the hygroscopic growth of the overall aerosol. Particle hygroscopicity may change as a function of time, place, and particle size (McMurry and Stolzenburg, 1989; Swietlicki *et al.*, 2008).

Estimation of hygroscopic growth factors with Köhler theory requires detailed knowledge of particle composition as well

as a thermodynamic model, which describes the concentration dependence of the water activity for such a mixture. The hygroscopic growth factor of a mixture, $gf_{\text{mix}}(\text{RH})$, can be calculated from the growth factors of the individual components of the aerosol and their respective volume fractions, V_k , by employing the Zdanovskii-Stokes-Robinson relation (ZSR relation) (Stokes and Robinson, 1966; Meyer *et al.*, 2009; Sjogren *et al.*, 2007; Stock *et al.*, 2011; Akpootu and Gana, 2013):

$$gf_{\text{mix}}(\text{RH}) = \left(\sum_k V_k gf_k^3 \right)^{1/3} \quad (2)$$

where the summation was performed over all compounds present in the particles. Solute-solute interactions are neglected in this model while the volume additivity was assumed. The model assumes spherical particles, ideal mixing (i.e. no volume change upon mixing) and independent water uptake of the organic and inorganic components.

This was also calculated using the corresponding number fractions n_k as (Meier *et al.*, 2009; Duplissy *et al.*, 2011; Akpootu and Gana, 2013).

$$gf_{\text{mix}}(\text{RH}) = \left(\sum_k n_k gf_k^3 \right)^{1/3} \quad (3)$$

where n_k is the number fraction of particles having the growth factor gf_k .

The $gf_{\text{mix}}(\text{RH})$ as a function of mass mix ratio has been proposed by Tijjani and Uba (2013) as reported by Akpootu and Abdul salami (2013) to be

$$gf_{\text{mix}}(\text{RH}) = \left(\sum_k m_k gf_k^3 \right)^{1/3} \quad (4)$$

The subscript k in the above equations represents the different substances.

The RH dependence of $gf_{\text{mix}}(\text{RH})$ was parameterized in a good approximation by a one-parameter equation (Petters and Kreidenweis, 2007; Akpootu and Abdul salami, 2013):

$$gf_{\text{mix}}(a_w) = \left(1 + \kappa \frac{a_w}{1-a_w} \right)^{1/3} \quad (5)$$

Here, a_w is the water activity, which can be replaced by the relative humidity RH at equilibrium (Seinfeld and Pandis,

2006), if the Kelvin effect is negligible, as for particles with sizes more relevant for light scattering and absorption. The coefficient κ is a simple measure of the particle's hygroscopicity and takes into consideration all solute properties (Raoult effect).

Humidograms of the ambient aerosols obtained in different atmospheric conditions revealed that $gf_{mix}(RH)$ could as well be fitted well with a γ -law (Swietlicki *et al.*, 2000; Gysel *et al.*, 2009; Putaud, 2012; Akpootu and Abdul salami, 2013) as

$$gf_{mix}(RH) = \left(1 - \frac{RH}{100}\right)^\gamma \quad (6)$$

Particle hygroscopicity is a measure that scales the volume of water associated with a unit volume of dry particle (Petters and Kreidenweis, 2007) and depends on the molar volume and the activity coefficients of the dissolved compounds (Christensen and Petters, 2012).

The bulk hygroscopicity factor under subsaturation RH conditions was determined using the following relation (Akpootu and Abdul salami, 2013):

$$B = (1 - gf_{mix}^3) \ln a_w \quad (7)$$

where a_w is the water activity that is replaced by the relative humidity as previously explained from equation (5).

RESULTS AND DISCUSSION

Table 2: The growth factor and bulk hygroscopicity of aerosols using number mix ratio for model 1-3

| RH(%) | Model 1 | | Model 2 | | Model 3 | |
|-------|------------|----------|------------|----------|------------|----------|
| | gf_{mix} | Bulk Hyg | gf_{mix} | Bulk Hyg | gf_{mix} | Bulk Hyg |
| 50 | 1.04355 | 0.09456 | 1.04061 | 0.08792 | 1.03798 | 0.08201 |
| 70 | 1.06846 | 0.07839 | 1.06392 | 0.07287 | 1.05987 | 0.06797 |
| 80 | 1.09380 | 0.06886 | 1.08770 | 0.06401 | 1.08224 | 0.05971 |
| 90 | 1.15139 | 0.05546 | 1.14195 | 0.05154 | 1.13349 | 0.04808 |
| 95 | 1.23233 | 0.04470 | 1.21864 | 0.04154 | 1.20631 | 0.03875 |
| 98 | 1.37277 | 0.03206 | 1.35260 | 0.02979 | 1.33430 | 0.02779 |
| 99 | 1.48942 | 0.02316 | 1.46448 | 0.02152 | 1.44177 | 0.02007 |

Table 2 shows that there is an overall increase in aerosol hygroscopic growth factor for number mix ratio model with increase in relative humidity from 50-99% RHs in each model. The bulk hygroscopicity decreases with increase in RH from 50 – 99% RHs for all the three models used.

More so, it was observed that the growth factor decreases with RHs from 50-99% RHs when the models were compared from model 1 to model 3. Similarly, the bulk hygroscopicity

decreases with RHs from model 1 to model 3. The aerosol growth factor revealed that the mixture is barely hygroscopic from 50 – 80% RHs, less hygroscopic from 90 – 95% RHs and more hygroscopic from 98 – 99% RHs for the number mix ratio. The bulk hygroscopicity ranges between 0.02316 to 0.09456 for model 1, 0.02152 to 0.08792 for model 2 and 0.02007 to 0.08201 for model 3.

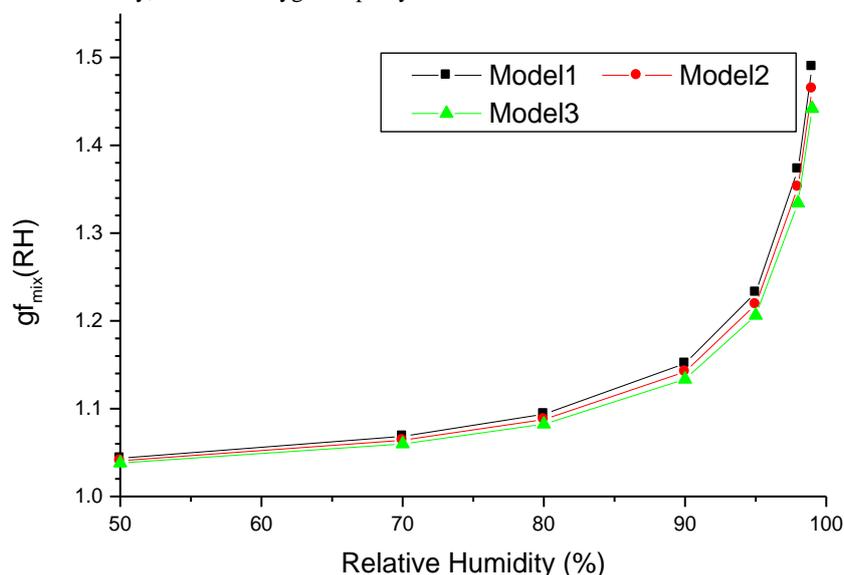


Figure 1: Growth factor of the mixture using number mix ratio (model 1-3)

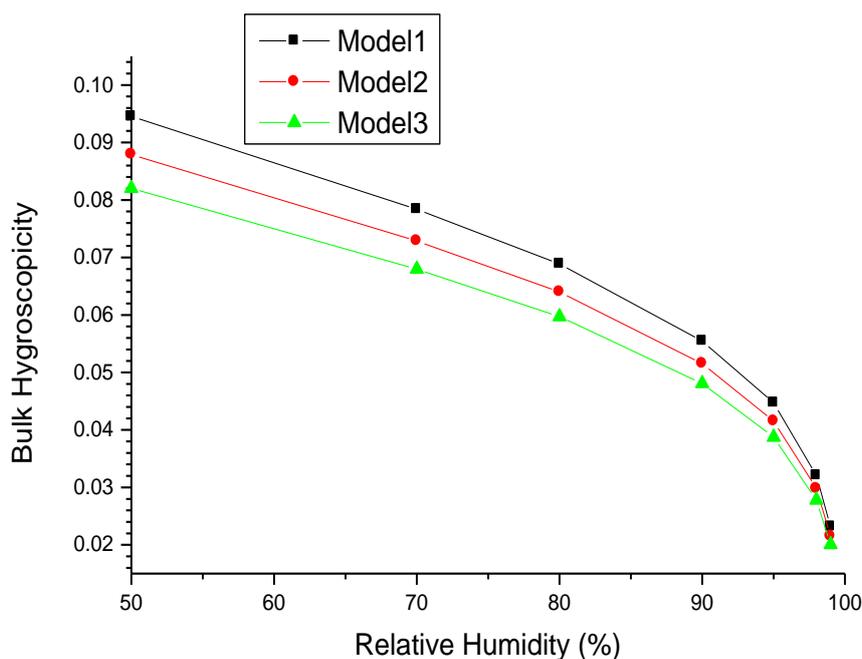


Figure 2: Bulk hygroscopicity of the mixture using number mix ratio (model 1-3)

Figure 1 depicts a non-linear increase in aerosol hygroscopic growth factor with RHs. The growth factor rise up steadily with increasing RH, this could be attributed to the fact that higher RH indicates more moisture content in the atmosphere which means the aerosol particles tends to absorb more water vapour in the atmosphere. The range of values estimated for

the gf_{mix} shown in table 2 the mixture as shown in figure 1 are described as barely hygroscopic, less hygroscopic and more hygroscopic growth in accordance with the description for the range of values by Swietlick et al. (2008), Liu et al. (2011). The bulk hygroscopicity decreases with increase in RHs as displayed in figure 2 for the three models.

Table 3: The growth factor for number mix ratio using model 1

| RH (%) | gf_{mix} | $RH/(1-RH)$ | gf^3 | $\ln(1-RH/100)$ | $\ln gf_{mix}$ |
|--------|------------|-------------|---------|-----------------|----------------|
| 50 | 1.04355 | 1.00000 | 1.13642 | -0.69315 | 0.04263 |
| 70 | 1.06846 | 2.33333 | 1.21978 | -1.20397 | 0.06622 |
| 80 | 1.09380 | 4.00000 | 1.30861 | -1.60944 | 0.08966 |
| 90 | 1.15139 | 9.00000 | 1.52639 | -2.30259 | 0.14097 |
| 95 | 1.23233 | 19.00000 | 1.87145 | -2.99573 | 0.20890 |
| 98 | 1.37277 | 49.00000 | 2.58699 | -3.91202 | 0.31683 |
| 99 | 1.48942 | 99.00000 | 3.30406 | -4.60517 | 0.39838 |

Table 4: The growth factor for number mix ratio using model 2

| RH (%) | gf_{mix} | $RH/(1-RH)$ | gf^3 | $\ln(1-RH/100)$ | $\ln gf_{mix}$ |
|--------|------------|-------------|---------|-----------------|----------------|
| 50 | 1.04061 | 1.00000 | 1.12684 | -0.69315 | 0.03981 |
| 70 | 1.06392 | 2.33333 | 1.20429 | -1.20397 | 0.06196 |
| 80 | 1.08770 | 4.00000 | 1.28683 | -1.60944 | 0.08406 |
| 90 | 1.14195 | 9.00000 | 1.48918 | -2.30259 | 0.13274 |
| 95 | 1.21864 | 19.00000 | 1.80978 | -2.99573 | 0.19774 |
| 98 | 1.35260 | 49.00000 | 2.47462 | -3.91202 | 0.30203 |
| 99 | 1.46448 | 99.00000 | 3.14086 | -4.60517 | 0.38150 |

Table 5: The growth factor for number mix ratio using model 3

| RH (%) | gf_{mix} | $RH/(1-RH)$ | gf^3 | $\ln(1-RH/100)$ | $\ln gf_{mix}$ |
|--------|------------|-------------|---------|-----------------|----------------|
| 50 | 1.03798 | 1.00000 | 1.11832 | -0.69315 | 0.03728 |
| 70 | 1.05987 | 2.33333 | 1.19057 | -1.20397 | 0.05814 |
| 80 | 1.08224 | 4.00000 | 1.26756 | -1.60944 | 0.07903 |
| 90 | 1.13349 | 9.00000 | 1.45631 | -2.30259 | 0.12530 |
| 95 | 1.20631 | 19.00000 | 1.75538 | -2.99573 | 0.18756 |
| 98 | 1.33430 | 49.00000 | 2.37555 | -3.91202 | 0.28841 |
| 99 | 1.44177 | 99.00000 | 2.99704 | -4.60517 | 0.36587 |

Tables 3, 4 and 5 shows the data estimated for the number mix ratio using equations (5) and (6). The results of the modeling using equations (12) and (13) are shown in table 6

Table 6: Summary of the results of R^2 , k , c and γ for the number mix ratio.

| Equations used | R^2 | k | constant | γ | Models used |
|----------------|---------|---------|----------|----------|-------------|
| 5 | 0.96171 | 0.02201 | 1.27405 | | Model 1 |
| 6 | 0.97935 | | -0.04843 | -0.09252 | |
| 5 | 0.96171 | 0.02045 | 1.25472 | | Model 2 |
| 6 | 0.97787 | | -0.04831 | -0.08879 | |
| 5 | 0.96171 | 0.01908 | 1.23761 | | Model 3 |
| 6 | 0.97645 | | -0.04804 | -0.08532 | |

The fitted curve can be represented by any of the empirical parameters in the form of either equation (5) or (6) However, it was observed that equation (6) gives a higher coefficient of determination, R^2 as compared to equation (5) for the three models indicating that the growth factor is well fitted with the γ -law as compared to the parameterization by one – parameter equation.

Table 7: The growth factor and bulk hygroscopicity of aerosols using volume mix ratio for model (1-3)

| RH (%) | Model 1 | | Model 2 | | Model 3 | |
|--------|------------|----------|------------|----------|------------|----------|
| | gf_{mix} | Bulk Hyg | gf_{mix} | Bulk Hyg | gf_{mix} | Bulk Hyg |
| 50 | 1.13844 | 0.32956 | 1.13706 | 0.32587 | 1.13571 | 0.32224 |
| 70 | 1.22855 | 0.30471 | 1.22672 | 0.30175 | 1.22491 | 0.29884 |
| 80 | 1.32001 | 0.29009 | 1.3179 | 0.28763 | 1.31578 | 0.28517 |
| 90 | 1.51813 | 0.26328 | 1.51575 | 0.26155 | 1.51341 | 0.25985 |
| 95 | 1.77055 | 0.23341 | 1.76819 | 0.23227 | 1.76590 | 0.23117 |
| 98 | 2.15557 | 0.18214 | 2.15357 | 0.18158 | 2.15164 | 0.18104 |
| 99 | 2.44351 | 0.13658 | 2.44184 | 0.13628 | 2.44008 | 0.13596 |

Table 7 shows that there is a general increase in aerosol hygroscopic growth factor for the volume mix ratio model with increase in RHs from 50-99% RHs in each model. However, the bulk hygroscopicity decreases with increase in RH from 50 – 99% RHs for all the three adopted models. It was observed that both the aerosol growth factor and bulk hygroscopicity decreases when compared from model 1 to model 3. The growth factor revealed that the mixture is less hygroscopic from 50 – 80% RHs, more hygroscopic from 90 – 95% RHs and most hygroscopic from 98 – 99% RHs for the volume mix ratio. The bulk hygroscopicity ranges between 0.13658 to 0.32956 for model 1, 0.13628 to 0.32587 for model 2 and 0.13596 to 0.32224 for model 3.

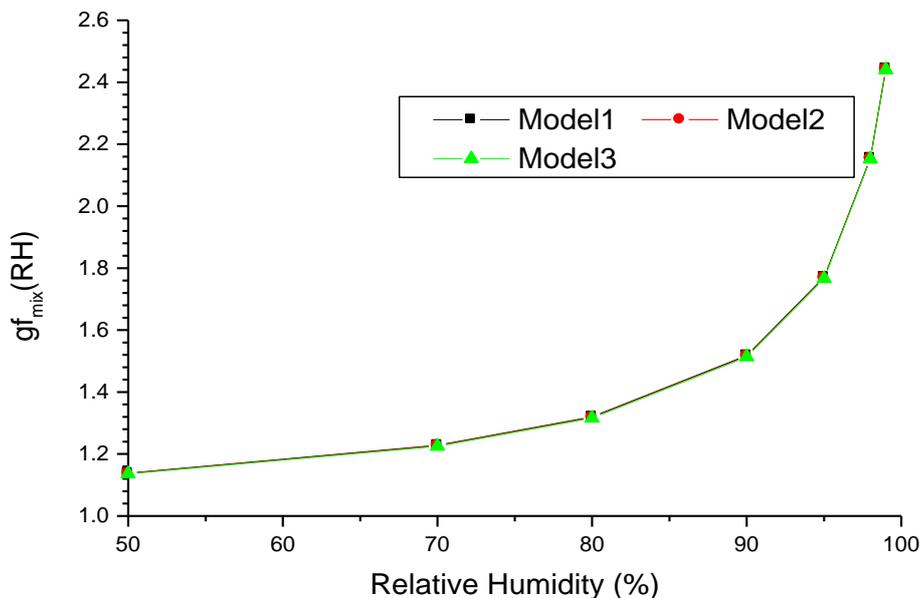


Figure 3: Growth factor of the mixture using volume mix ratio (model 1-3)

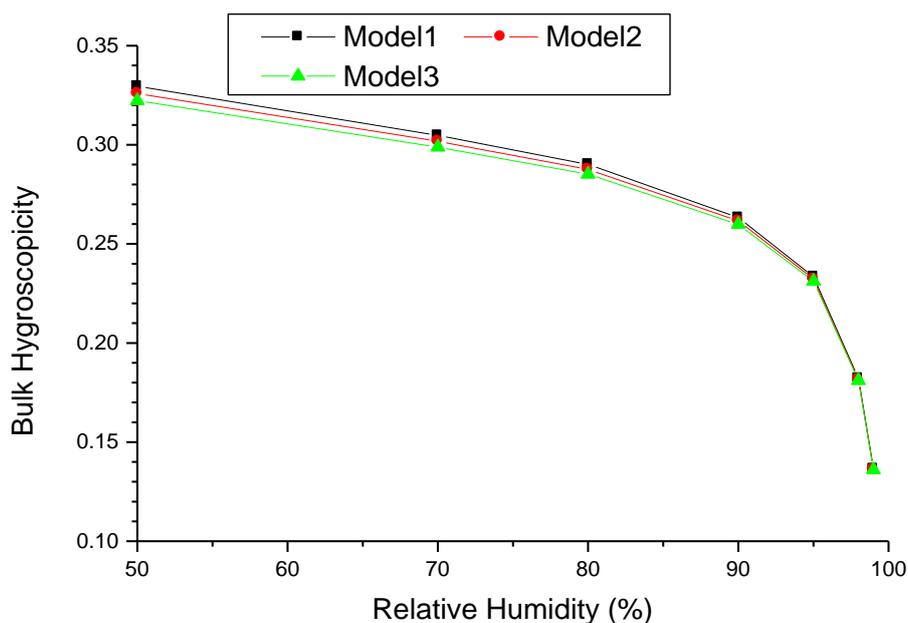


Figure 4: Bulk hygroscopicity of the mixture using volume mix ratio (model 1-3)

Figure 3 depicts a non-linear increase in aerosol hygroscopic growth factor with RHs, however, the rate of increase appears to be almost constant. The growth factor rises up steadily with increasing RH, this may be attributed to the fact that higher RH implies presence of more atmospheric moisture contents making the aerosol particles to absorb more water vapour. The range of values estimated for the gf_{mix} shown in table 7 the

mixture as depicted in figure 3 are described as less hygroscopic, more hygroscopic and most hygroscopic growth in accordance with the description for the range of values by Swietlick et al. (2008), Liu et al. (2011). The bulk hygroscopicity decreases with increase in RHs as displayed in figure 4 for the three models.

Table 8: The growth factor for volume mix ratio using model 1

| RH (%) | gf_{mix} | $RH/(1-RH)$ | gf^3 | $\ln(1-RH/100)$ | $\ln gf_{mix}$ |
|--------|------------|-------------|----------|-----------------|----------------|
| 50 | 1.13844 | 1.00000 | 1.47546 | -0.69315 | 0.12966 |
| 70 | 1.22855 | 2.33333 | 1.85430 | -1.20397 | 0.20584 |
| 80 | 1.32001 | 4.00000 | 2.30001 | -1.60944 | 0.27764 |
| 90 | 1.51813 | 9.00000 | 3.49884 | -2.30259 | 0.41748 |
| 95 | 1.77055 | 19.00000 | 5.55042 | -2.99573 | 0.57129 |
| 98 | 2.15557 | 49.00000 | 10.01579 | -3.91202 | 0.76805 |
| 99 | 2.44351 | 99.00000 | 14.58963 | -4.60517 | 0.89344 |

Table 9: The growth factor for volume mix ratio using model 2

| RH (%) | gf_{mix} | $RH/(1-RH)$ | gf^3 | $\ln(1-RH/100)$ | $\ln gf_{mix}$ |
|--------|------------|-------------|----------|-----------------|----------------|
| 50 | 1.13706 | 1.00000 | 1.47013 | -0.69315 | 0.12845 |
| 70 | 1.22672 | 2.33333 | 1.84601 | -1.20397 | 0.20434 |
| 80 | 1.31790 | 4.00000 | 2.28900 | -1.60944 | 0.27604 |
| 90 | 1.51575 | 9.00000 | 3.48241 | -2.30259 | 0.41591 |
| 95 | 1.76819 | 19.00000 | 5.52828 | -2.99573 | 0.56996 |
| 98 | 2.15357 | 49.00000 | 9.98793 | -3.91202 | 0.76713 |
| 99 | 2.44184 | 99.00000 | 14.55972 | -4.60517 | 0.89275 |

Table 10: The growth factor for volume mix ratio using model 3

| RH (%) | gf_{mix} | $RH/(1-RH)$ | gf^3 | $\ln(1-RH/100)$ | $\ln gf_{mix}$ |
|--------|------------|-------------|----------|-----------------|----------------|
| 50 | 1.13571 | 1.00000 | 1.46489 | -0.69315 | 0.12726 |
| 70 | 1.22491 | 2.33333 | 1.83786 | -1.20397 | 0.20287 |
| 80 | 1.31578 | 4.00000 | 2.27796 | -1.60944 | 0.27443 |
| 90 | 1.51341 | 9.00000 | 3.46634 | -2.30259 | 0.41437 |
| 95 | 1.76590 | 19.00000 | 5.50676 | -2.99573 | 0.56866 |
| 98 | 2.15164 | 49.00000 | 9.96119 | -3.91202 | 0.76623 |
| 99 | 2.44008 | 99.00000 | 14.52822 | -4.60517 | 0.89203 |

Tables 8, 9 and 10 shows the data obtained for the volume mix ratio using equations (5) and (6). The results of the modeling using equations (5) and (6) are shown in table 11.

Table 11: Summary of the results of R^2 , k , c and γ for the volume mix ratio.

| Equations used | R^2 | k | constant | γ | Models used |
|----------------|---------|---------|----------|----------|-------------|
| 5 | 0.97083 | 0.13455 | 2.08813 | | Model 1 |
| 6 | 0.99787 | | -0.0311 | -0.201 | |
| 5 | 0.97105 | 0.13432 | 2.07683 | | Model 2 |
| 6 | 0.99781 | | -0.0328 | -0.2012 | |
| 5 | 0.97123 | 0.13408 | 2.06598 | | Model 3 |
| 6 | 0.99775 | | -0.0345 | -0.2013 | |

The fitted curve can be represented by any of the empirical parameters in the form of either equation (5) or (6) However, it was observed that equation (6) gives a higher coefficient of determination, R^2 as compared to equation (5) for the three

models indicating that the growth factor is well fitted with the γ -law as compared to the parameterization by one – parameter equation.

Table 12: The growth factor and bulk hygroscopicity of aerosols using mass mix ratio for model (1-3)

| RH (%) | Model 1 | | Model 2 | | Model 3 | |
|--------|-------------------|----------|-------------------|----------|-------------------|----------|
| | gf _{mix} | Bulk Hyg | gf _{mix} | Bulk Hyg | gf _{mix} | Bulk Hyg |
| 50 | 1.12707 | 0.29925 | 1.12625 | 0.29706 | 1.12542 | 0.29487 |
| 70 | 1.20785 | 0.27183 | 1.20670 | 0.27004 | 1.20554 | 0.26823 |
| 80 | 1.29124 | 0.25726 | 1.28984 | 0.25569 | 1.28847 | 0.25418 |
| 90 | 1.47660 | 0.23385 | 1.47494 | 0.23270 | 1.47325 | 0.23154 |
| 95 | 1.72237 | 0.21079 | 1.72052 | 0.20995 | 1.71873 | 0.20913 |
| 98 | 2.10781 | 0.16899 | 2.10610 | 0.16853 | 2.10448 | 0.16809 |
| 99 | 2.39975 | 0.12884 | 2.39819 | 0.12857 | 2.39671 | 0.12831 |

Table 12 shows that there is a general increase in aerosol hygroscopic growth factor for the mass mix ratio model with increase in RHs from 50-99% RHs in each model. However, the bulk hygroscopicity decreases with increase in RH from 50 – 99% RH for all the three models.

It was observed that both the aerosol growth factor and bulk hygroscopicity decreases when compared from model 1 to

model 3. The growth factor revealed that the mixture is less hygroscopic from 50 – 80% RHs, more hygroscopic from 90 – 95% RHs and most hygroscopic from 98 – 99% RHs for the mass mix ratio. The bulk hygroscopicity ranges between 0.12884 to 0.29925 for model 1, 0.12857 to 0.29706 for model 2 and 0.12831 to 0.29487 for model 3.

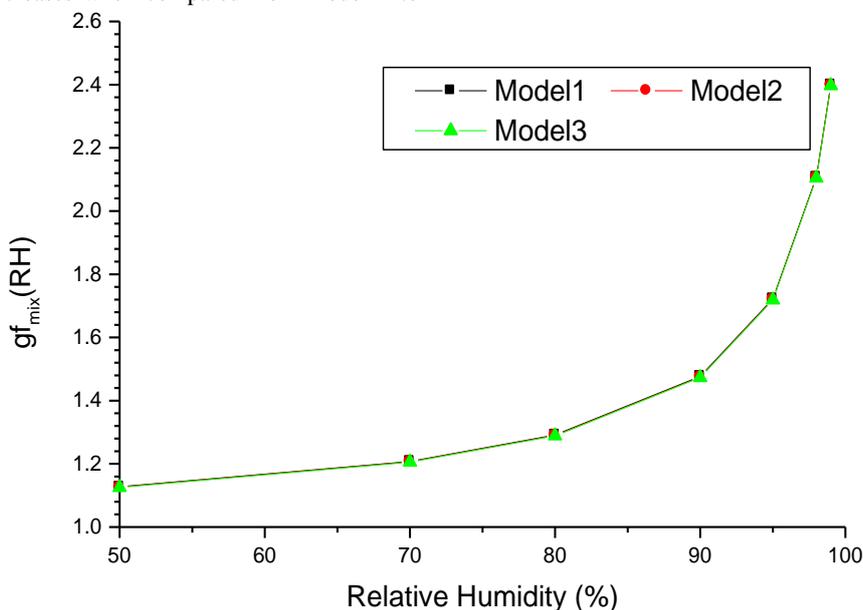


Figure 5: Growth factor of the mixture using mass mix ratio (model 1-3)

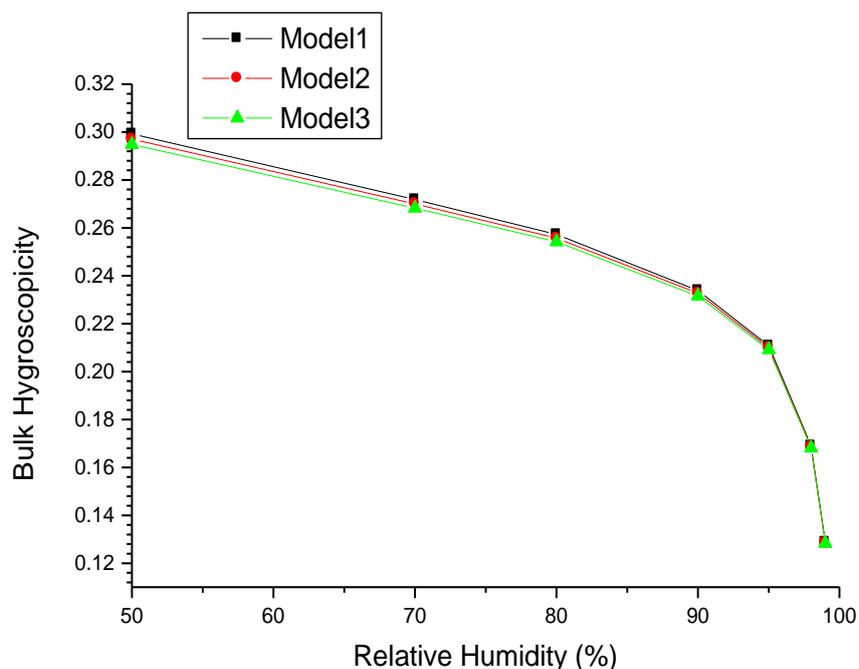


Figure 6: Bulk hygroscopicity of the mixture using mass mix ratio (model 1-3)

Figure 5 depicts a non-linear increase in aerosol hygroscopic growth factor with RHs, however, the rate of increase appears to be almost constant. The growth factor rises up steadily with increasing RH, this may be attributed to the fact that higher RH implies more atmospheric moisture content which makes the aerosol particles to absorb more water vapour on particle surface. The range of values estimated for the gf_{mix} shown in

table 12 the mixture as depicted in figure 5 are described as less hygroscopic, more hygroscopic and most hygroscopic growth in accordance with the description for the range of values by Swietlick et al. (2008), Liu et al. (2011). The bulk hygroscopicity decreases with increase in RHs as displayed in figure 6 with almost constant rate for the three models under study.

Table 13: The growth factor for mass mix ratio using model 1

| RH (%) | gf_{mix} | $RH/(1-RH)$ | gf^3 | $\ln(1-RH/100)$ | $\ln gf_{mix}$ |
|--------|------------|-------------|----------|-----------------|----------------|
| 50 | 1.12707 | 1.00000 | 1.43172 | -0.69315 | 0.11963 |
| 70 | 1.20785 | 2.33333 | 1.76212 | -1.20397 | 0.18884 |
| 80 | 1.29124 | 4.00000 | 2.15289 | -1.60944 | 0.25560 |
| 90 | 1.47660 | 9.00000 | 3.21949 | -2.30259 | 0.38974 |
| 95 | 1.72237 | 19.00000 | 5.10955 | -2.99573 | 0.54370 |
| 98 | 2.10781 | 49.00000 | 9.36473 | -3.91202 | 0.74565 |
| 99 | 2.39975 | 99.00000 | 13.81964 | -4.60517 | 0.87536 |

Table 14: The growth factor for mass mix ratio using model 2

| RH (%) | gf_{mix} | $RH/(1-RH)$ | gf^3 | $\ln(1-RH/100)$ | $\ln gf_{mix}$ |
|--------|------------|-------------|----------|-----------------|----------------|
| 50 | 1.12625 | 1.00000 | 1.42856 | -0.69315 | 0.11889 |
| 70 | 1.20670 | 2.33333 | 1.75711 | -1.20397 | 0.18789 |
| 80 | 1.28984 | 4.00000 | 2.14587 | -1.60944 | 0.25451 |
| 90 | 1.47494 | 9.00000 | 3.20863 | -2.30259 | 0.38861 |
| 95 | 1.72052 | 19.00000 | 5.09304 | -2.99573 | 0.54262 |
| 98 | 2.10610 | 49.00000 | 9.34193 | -3.91202 | 0.74484 |
| 99 | 2.39819 | 99.00000 | 13.79267 | -4.60517 | 0.87471 |

Table 15: The growth factor for mass mix ratio using model 3

| RH (%) | gf_{mix} | $RH/(1-RH)$ | gf^3 | $\ln(1-RH/100)$ | $\ln gf_{mix}$ |
|--------|------------|-------------|----------|-----------------|----------------|
| 50 | 1.12542 | 1.00000 | 1.42541 | -0.69315 | 0.11815 |
| 70 | 1.20554 | 2.33333 | 1.75203 | -1.20397 | 0.18692 |
| 80 | 1.28847 | 4.00000 | 2.13908 | -1.60944 | 0.25346 |
| 90 | 1.47325 | 9.00000 | 3.19761 | -2.30259 | 0.38747 |
| 95 | 1.71873 | 19.00000 | 5.07716 | -2.99573 | 0.54158 |
| 98 | 2.10448 | 49.00000 | 9.32035 | -3.91202 | 0.74407 |
| 99 | 2.39671 | 99.00000 | 13.76718 | -4.60517 | 0.87410 |

Tables 13, 14 and 15 shows the data obtained for the mass mix ratio using equations (5) and (6). The results of the modeling using equations (5) and (6) are shown in table 16.

Table 16: Summary of the results of R^2 , k , c and γ for the mass mix ratio.

| Equations used | R^2 | k | constant | γ | Models used |
|----------------|---------|---------|----------|----------|-------------|
| 5 | 0.97531 | 0.12738 | 1.92953 | | Model 1 |
| 6 | 0.99622 | | -0.04749 | -0.19923 | |
| 5 | 0.97546 | 0.12715 | 1.92241 | | Model 2 |
| 6 | 0.99615 | | -0.04852 | -0.19927 | |
| 5 | 0.97560 | 0.12694 | 1.91534 | | Model 3 |
| 6 | 0.99608 | | -0.04956 | -0.19932 | |

The fitted curve can be represented by any of the empirical parameters in the form of either equations (5) or (6) However, it was observed that equation (6) gives a higher coefficient of determination, R^2 as compared to equation (5) for the three models indicating that the growth factor is well fitted with the γ -law as compared to the parameterization by one – parameter equation.

CONCLUSION

The analysis in this study shows that the aerosol hygroscopic growth factor gf_{mix} increases with increase in RH while the bulk hygroscopicity factor decreases with increase in RH. The growth factor indicates that the mixture is barely hygroscopic, less hygroscopic, more hygroscopic for the number mix ratio and it's less hygroscopic, more hygroscopic and most hygroscopic for the volume and mass mix ratios. The bulk hygroscopicity ranges between 0.02007 to 0.09456 for the number mix ratio from model 1 to model 3, the bulk hygroscopicity ranges between 0.13596 to 0.32956 for the volume mix ratio from model 1 to model 3 while the bulk hygroscopicity ranges between 0.12831 to 0.29925 for the mass mix ratio from model 1 to model 3. The growth factor is well fitted with the γ -law as compared to the parameterization by one – parameter equation based on the coefficient of determination. The number mix, volume mix and mass mix ratios shows an increase in particle diameter with increase in RH with a steep curve of deliquescence found from 95-99% RHs. However, the volume mix ratio shows more increase in gf_{mix} with RHs and gives higher coefficient of determination when compared to the number mix ratio and mass mix ratio. The results showed that the coefficient of determination, $R^2 > 96\%$ for all the three models used in this study.

REFERENCES

Akande, J. O., Momoh, M., Saidu, I. G., Iliyasu, M. I., Akpootu, D. O., Abubakar, M. B and Abdullahi, M. B (2013). Evaluation of Residence Time of Dust Aerosols during the Harmattan Season in Sokoto Area of Northwestern Nigeria, Using Visibility Data. *IOSR Journal of Environmental Science, Toxicology And Food Technology.*, **4**(4), 71-74.

Akpootu, D. O and Abdul salami, M. J (2013). The Optical and Microphysical Properties of water Soluble Aerosols. *The International Journal of Engineering And Science.*, **2**(12), 66-79.

Akpootu, D. O and Gana, N. N (2013). The Effect of Relative Humidity on the Hygroscopic Growth Factor and Bulk Hygroscopicity of water Soluble Aerosols. *The International Journal Of Engineering And Science.*, **2**(11), 48-57.

Akpootu, D. O and Momoh, M (2013a). The Ångström Exponent and Turbidity of Soot Component in the Radiative Forcing of Urban Aerosols. *Nigerian Journal of Basic and*

Applied Science., **21**(1), 70-78. ISSN: 0794-5698: DOI: <http://dx.doi.org/10.4314/njbas.v21i1.11>

Akpootu, D. O and Momoh, M (2013b). The Scattering Coefficient, Extinction Coefficient and Single Scattering albedo of water Soluble in the Radiative Forcing of Urban Aerosols. *Scholars Research Library.*, **5**(2), 109-120. ISSN: 0975-508X

Akpootu, D. O and Muhammad, S. B (2013). The Angstrom Exponent and Turbidity of Soot Component in the Radiative Forcing of Urban Aerosols. A paper presented at the 35th Annual Nigerian Institute of Physics National Conference held at the University of Abuja, Nigeria on the 25th-30th March, 2013.

Akpootu, D. O and Sharafa, S. B. (2013). The scattering coefficient, extinction coefficient and single scattering albedo of soot in the radiative forcing of urban aerosols, *Scholars Research Library.*, **4**(3), 31-41.

Akpootu, D. O and Tijjani, B. I (2014). The Aerosol Hygroscopic Growth Factor and Bulk Hygroscopicity of Soot in the Hygroscopicity of Urban Aerosols. A paper presented at the 24th Annual Congress and Colloquium of the Nigerian Association of Mathematical Physics held at University of Benin, Benin City, Nigeria on the 25th-28th February, 2014.

Bond, T.C. and Bergstrom, R.W. (2006): Light absorption by Carbonaceous particles: An Investigative Review, *Aerosol Sci. Technol.*, **40**, 27-67.

Christensen, S. I. and Petters, M. D. (2012): The role of temperature in cloud droplet activation, *J. Phys. Chem. A* **116**(39): 9706–9717.

Duplissy, J., DeCarlo, P. F., Dommen, J., Alfarra, M. R., Metzger, A., Barmadimos, I., Prevot, A. S. H., Weingartner, E., Tritscher, T., Gysel, M., Aiken, A. C., Jimenez, J. L., Canagaratna, M. R., Worsnop, D. R., Collins, D. R., Tomlinson, J., and Baltensperger, U. (2011): Relating hygroscopicity and composition of organic aerosol particulate matter, *Atmos. Chem. Phys.*, **11**, 1155– 1165, doi:10.5194/acp-11-1155-2011.

Essienimo, A. O., Momoh, M and Akpootu, D. O. (2015a). Differential Particle Size Distribution of Aerosol across North Western Region of Nigeria. *International Research Journal of Engineering and Technology (IRJET).*, **2**(9), 555-561. e-ISSN: 2395 -0056, p-ISSN: 2395-0072.

Essienimo, A. O., Momoh, M and Akpootu, D. O. (2015b). Percentage Composition of Particle Size Distribution of Aerosol mass concentration during 2014 Winter Season for

- some Selected Regions in Northwestern Nigeria. *International Research Journal of Engineering and Technology (IRJET)*, 2(9), 562-569. e-ISSN: 2395 -0056, p-ISSN: 2395-0072.
- Essienimo, A. O., Momoh, M and Akpootu, D. O. (2016a). Mass Concentration gradient of Aerosol Across Selected States in Northwestern Nigeria. *International Journal of Technology Enhancements and Emerging Engineering Research*, 4(1), 17 – 22. ISSN: 2347-4289
- Essienimo, A. O., Momoh, M and Akpootu, D. O. (2016b). Seasonal and Monthly Estimation of Mean Residence Time of the Harmattan Dust in Kano, Northern Nigeria Using Horizontal Visibility Data. *International Journal of Technology Enhancements and Emerging Engineering Research*, 4(2), 9 – 15. ISSN 2347-4289
- Gysel, M., McFiggans, G. B., and Coe, H. (2009): Inversion of tandem differential mobility analyser (tdma) measurements, *J. Aerosol Sci.*, 40, 134–151, doi:10.1016/j.jaerosci.2008.07.013.
- Hess, M., Koepke, P and Schult, I (1998): Optical Properties of Aerosols and Clouds. *American Meteorology Society*.
- Intergovernmental Panel on Climate Change IPCC. (2013): Summary for Policymakers, in: Climate Change: The Physical Science Basis. Contribution of Working Group I to the sixth Assessment Report Cambridge University Press, Cambridge, UK and New York, NY, USA.
- IPCC (2007): Climate Change 2007: The Scientific Basis. In Solomon, S., Ding, Y., Griggs, D.G., Noguera, M., Vanderlinden, P.G., Dai, X., Maskell, K. and Johnson, C.A. (Eds). Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change. Cambridge University Press. Cambridge.
- Liu, P. F., Zhao, C. S., Gobel, T., Hallbauer, E., Nowak, A., Ran, L., Xu, W. Y., Deng, Z. Z., Ma, N., Milderger, K., Henning, S., Stratmann, F., and Wiedensohler, A. (2011): Hygroscopic Properties of Aerosol Particles at High Relative Humidity and their Diurnal Variations in the North China Plain, *Atm. Chem. Phys.*, Discuss., 11, 2991-3040.
- Lohmann, U. and Feichter, J. (2005) Global indirect aerosol effects: a review. *Atmospheric Chemistry and Physics* 5(3), 715–737.
- McFiggans, G., Artaxo, P., Baltensperger, U., Coe, H and Facchini, M.C (2006): The effect of physical and chemical aerosol properties on warm cloud droplet activation *Atmos. Chem. Phys.* 6, 2593-2649.
- McMurry, P.H., and Stolzenburg, M.R. (1989): On the sensitivity of particle size to relative humidity for Los Angeles aerosols. *Atmos Environ.*, 23:497–507. doi:10.1016/0004-6981(89) 90593-3
- Meier J., Wehner, B., Massling, A., Birmili, W., Nowak, A., Gnauk, T., Brüggemann, E., Herrmann, H., Min, H., and Wiedensohler, A. (2009): Hygroscopic growth of urban aerosols particles in Beijing (China) during wintertime: A comparison of three experimental methods., *Atmos. Chem. Phys.*, 9, 6865–6880, www.atmos-chemphys.net/9/6865/2009/
- Meseke, N. O., Akpootu, D. O., Falaiye, O. A. and Targema, T. V. (2022). COMPARATIVE ASSESSMENT OF PARTICULATE MATTER USING LOW COST SENSOR: A CASE STUDY OF ABUJA AND KANO, NIGERIA. *FUDMA Journal of Sciences (FJS)*, Vol. 6 No. 3, June, 2022, pp 290 – 300. ISSN 2645 – 2944. DOI: https://doi.org/10.33003/fjs-2022-0604-1066
- Meyer, N. K., Duplissy, J., Gysel, M., Metzger, A., Dommen, J., Weingartner, E., Alfarra, M. R., Prevot, A. S. H., Fletcher, C., Good, N., McFiggans, G., Jonsson, A. M., Hallquist, M., Baltensperger, U., and Ristovski, Z. D. (2009): Analysis of the hygroscopic and volatile properties of ammonium sulphate seeded and unseeded SOA particles, *Atmos. Chem. Phys.*, 9, 721–732, doi:10.5194/acp-9-721-2009.
- Pandis, S. N., Wexler, A. S., and Seinfeld, J. H. (1995): Dynamics of tropospheric aerosols, *J. Phys. Chem.*, 99, 9646–9659.
- Peters, M. D. and Kreidenweis, S. M. (2007): A single parameter representation of hygroscopic growth and cloud condensation nucleus activity, *Atmos. Chem. Phys.*, 7, 1961–1971, doi:10.5194/acp-7-1961-2007.
- Putaud, J.P. (2012): Aerosol hygroscopicity at Ispra EMEP-GAW station by M. Adam et al., *Atmos. Chem. Phys. Discuss.*, 12, C1316–C1322
- Randles, C. A., Russell, L. M., and Ramaswamy, V. (2004): Hygroscopic and optical properties of organic sea salt aerosol and consequences for climate forcing, *Geographical Research Letters*, Vol. 31, LI6108, doi: 10, 1029/2004GL020628.
- Satheesh, S.K and Krishna Moorthy, K. (2005): Radiative effects of natural aerosols: A Review. *Atmos. Environ* 39, 2089-2110.
- Seinfeld, J. H. and Pandis, S. N. (2006): Atmospheric Chemistry and Physics: From Air Pollution to Climate Change, John Wiley and Sons, Inc., New York.
- Seinfeld, J. H., Bretherton, C., Carslaw, K. S and Wood, R. (2016). Improving our fundamental understanding of the role of aerosol- cloud interactions in the climate system. *Proceedings of the National Academy of Sciences* 113(21), 5781–5790.
- Sjogren, S., Gysel, M., Weingartner, E., Baltensperger, U., Cubison, M. J., Coe, H., Zardini, A. A., Marcolli, C., Krieger, U. K., and Peter, T. (2007): Hygroscopic growth and Water uptake kinetics of two-phase aerosol particles consisting of ammonium sulfate, adipic and humic acid mixtures, *J. Aerosol Sci.*, 38, 157–171, doi:10.1016/j.jaerosci.2006.11.005.
- Sloane, C. S. and Wolff, G. T. (1985): Prediction of ambient light-scattering using a physical model responsive to relative-humidity – validation with measurements from detroit, *Atmos. Environ.*, 19, 669–680.
- Stock, M., Cheng, Y. F. Birmili, W., Massling, A., Wehner, B., Müller, T., Leinert, S., Kalivitis, N., Mihalopoulos, N., and Wiedensohler, A. (2011): *Atmos. Chem. Phys.*, 11, 4251–4271, www.atmos-chemphys.net/11/4251/2011/ doi:10.5194/acp-11-4251-2011

Stokes, R. H., and Robinson, R. A. (1966): Interactions in aqueous nonelectrolyte solutions. I. Solute solvent equilibria, *J. Phys. Chem.*, 70, 2126–2130.

Swietlicki, E., Hansson, H. C., Hämeri, K., Svenningsson, B., Massling, A., McFiggans, G., McMurry, P. H., Petäjä, T., Tunved, P., Gysel, M., Topping, D., Weingartner, E., Baltensperger, U., Rissler, J., Wiedensohler, A., and Kulmala, M. (2008): Hygroscopic properties of submicrometer atmospheric aerosol particles measured with H-TDMA instruments in various environments 10 – a review, *Tellus B*, 60(3), 432–469, 6892, 6898, 6906, 6909

Swietlicki, E., Zhou, J., Covert, D. S., Hameri, K., Busch, B., Vakeva, M., Dusek, U., Berg, O. H., Wiedensohler, A., Aalto, P., Makela, J., Marinsson, B. G., Papaspiropoulos, G., Mentes, B., Frank, G. and Strach. F. (2000): Hygroscopic properties of aerosol particles in the north-eastern Atlantic during ACE-2. *Tellus* 52B, 201–227.

Tijjani, B. I and Akpootu, D. O (2012). The Effect of Soot and water Soluble in Radiative Forcing of Urban Aerosols. *International Journal of Research and Reviews in Pharmacy and Applied Science.*, 2(6), 1128-1143.

Tijjani, B. I and Akpootu, D. O (2013a). The Effect of Water Solubles in Radiative Forcing of Urban Aerosols. *Journal of the Nigerian Association of Mathematical Physics.*, Volume 23, 313-324. ISSN: 1116-4336

Tijjani, B. I and Akpootu D. O (2013b). The Effect of Soot in the Radiative Forcing of Urban Aerosols. *Journal of the Nigerian Association of Mathematical Physics.*, Volume 23, 325-334. ISSN: 1116-4336

Tijjani, B. I and Uba, S. (2013): The effect of hygroscopic growth on desert aerosols, *Pelagia Research Library.*, 4(4), 165-178.



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