



## ASSESSMENT OF PRODUCTIVITY STATUS OF VERTISOLS IN SUDAN SAVANNA ZONE OF NIGERIA

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### ABSTRACT

Higher water and nutrient retention of vertisols make it an important soil order with greater potentials for intensified agricultural production, especially in Sahel and Sudan Savannah; where rainfall is inadequate and erratic, as well as poor soil nutrient base. Hence, its utilization begs for research intervention for sustainable management. One hundred and twenty (120) each of composite and core samples were collected from the eastern part of Jigawa state (Longitude 12. 14 – 12. 69°N; Latitude 9. 54 – 10. 21° E) at three varying depths: 0 – 30, 30 – 60 and 60 – 90 cm. The samples were prepared and analysed employing standard procedures. Result after statistically analysis showed that, the Vertisols pH status ranged from moderately acidic at the surface (0 – 30 cm) through slightly acidic (30 – 60 cm) and moderately alkaline (60 – 90 cm). Organic carbon, total nitrogen and available phosphorus were low in all the depths. Most of the exchangeable bases were medium to high and a medium cation exchange capacity throughout. The Vertisols was texturally sandy clay loam throughout, highly aggregated and low in permeability. It can be inferred that, Vertisols in the Sudan Savannah Zone of Nigeria has good production potential to rice and Wheat crops. But maize and tomato could not thrive well due to little permeability observed, which indicates waterlogging tendency. However, complimentary application of organic and mineral fertilizers is required to argument its nutrient reserve for optimum and sustained production. Incorporation of crop residue and off season ploughing will go a long way in enhancing its water hydraulic conductivity (permeability) and reduction of surface caking which facilitates runoff and nutrient loss.

**Keywords:** Vertisols, productivity, water and nutrient retention; Sudan Savannah, sustainability.

### INTRODUCTION

Vertisols is one of the twelve soil orders (Soil survey staff, 2000, Michael and Dunn, 2000, Brady and Weil 1999) and is characteristically dark in colour, hence the common names; Black Earths, Black Cotton soils, Dark Cracking soils (Dudal 1965). However, that does not signify organic matter content. The name 'Vertisols' was derived from a Latin word *Vetere/Verito*, meaning to turn (Brady and Weil, 1999; Jutzi, *et al.*, 1988). Inconsistencies exist in the description of Vertisols between the classification systems used. United State Soil Taxonomy describes Vertisols as mineral soils that have a mesic, isomesic or warmer soil temperature regime; that do not have a lithic or paralithic contact or petrocalcic horizon or duripan within 50 cm of the soil surface; that, after the upper soil to a depth of 18 cm has been mixed, have 30% or more clay in all horizons down to a depth of 50 cm or more; that at some period in most years have cracks that are open to the surface or to the base of a lower layer and are at least 1cm wide at a depth of 50 cm unless the soil is irrigated; and that have one or more of the following characteristics: Gilgai; at some depth between

25 cm and 1m, slickensides close enough to intersect; or at some depth between 25 cm and 1 m, wedge – shaped (sphenoid), structural aggregates whose long axes are tilted 10° to 60° from the horizontal."

The World Reference Base (WRB) system, on the other hand describes Vertisols as soils with high shrink – swell properties having diagnostic ochric epepedon, usually without argillic epepedon (Michael and Dunn, 2000). However, Pal *et al.* (2009) had reported that Vertisols can also have argillic horizon. Vertisols are soils with high content of clay minerals that shrink and swell as they change water content. The shrinking – expansion property is as a result of the presence of interparticle and intraparticle porosity with changes with moisture content (Pal *et al.*, 2012). This is in contrast to the commonly invoked process of clay interlayer hydration – dehydration to explain shrink – swell phenomenon (Pal *et al.*, 2012). As Vertisols dry, the contraction property brings about wide cracks, which permit deep water percolation at initial entry and subsequently becomes very slow upon swelling. Vertisols have a more uniform soil profile; a self – mixed.

### Distribution of Vertisols

In the taxonomic nomenclature, Vertisols bear formative element 'ert' and are most commonly formed in warm, sub – humid or semi – arid climates (Jutzi, et al., 1988) where the natural vegetation is predominantly grass, open forest, or desert shrub (Pal, et al., 2012). Globally, Vertisols comprise of about 2.5% (about 250 mha) of the total land area (Brady and Weil, 1999). Large areas of Vertisols coverage are found in Northeastern Africa, India, and Australia with smaller areas scattered worldwide (Smith, 2016; Ahmad, 1996). Five countries; namely; Australia, India, Sudan, Chad and Ethiopia contain over 80 % of the total area of Vertisols coverage in the world (Virmani, et al., 2000; Wilding and Coulombe, 1996). In Africa, Sudan, Chad and Ethiopia are the countries with larger area of Vertisols: 40 mha, 16.5 mha and 10 mha respectively (Asnakew, 1999). Vertisols in Nigeria occur over an area of about 4 mha between latitudes 8°30' and 12°30'N and longitudes 10° and 14° E within Savannah ecological zone of the country (Klinkenberg and Higgins, 1968). Ojanuga, (2006); Kowal and Knabe, (1972) reported that in Nigeria, Vertisols predominantly occur in the Sudan and Northern Guinea Savannah Zones in the northwest and northeast of the country. Generally, Vertisols are most commonly found in lower landscape such as depression, bottomlands, lower foot slope (Pal, et al., 2012). In Jigawa state, even though there is no inventory of area of coverage of Vertisols, but its large fraction is found in the eastern part of the state within Hadejia river basin. The area is topographically low and flooded for part of every year.

Although Vertisols cover only a small area of the world's land surface and only a sub-dominant portion of any geographical zone, but is an important soil order particularly in drylands (Adeoye, 1992). The major factor contributing to the productivity of Vertisols in semi – arid environments is its high water-holding capacity (Virmani, et al., 2000). In this area, rainfall is uncertain, inadequate and highly variable (Ojanuga, 2006). For this, the ability of Vertisols to store sufficient water to carry crops through droughty periods is of great importance. More so, high surface area that ranged from 628 to 770 m<sup>2</sup>/g (Abdu and Uyovbisere, 2007, Adeoye, 1992) could support higher production by increasing nutrient retention, which is a limiting factor in Sudan savannah soils (Ojanuga, 2006). Because of its high clay content, vertisols is as well resilient to degradation compared to other soils. As such, Vertisols is considered significant global resources that serve as lifeline in subsistence agriculture. However, some characteristics, viz: lesser water intake, more hazards of waterlogging, difficult tillage (due to hardness and stickiness when dry and when wet respectively), poor seedbed conditions, and surface crusting, (Pal, et al., 2012; Dudal, 1973) do pose constraints for cultivation of diverse crops as well as sustained utilization. And some of these problems assume greater importance where a number of crop choices are needed for sustainable intensification to meet huge demand for food and product by man and industries (Myers and Pathak, 2001; Syers, et al., 2001;

Virmani et al., 2000). Efforts toward comprehensive and successful utilization of Vertisols are imperative for continued productivity and long – term sustainability.

Though Vertisols make up a relatively homogeneous soil order in taxonomic sense, but it is worth notice that they show diverse characteristics that are very important in agronomic context for their suitability to various crop productions. Therefore, adopting a general management and utilization recommendations may be misleading and detrimental. Based on the foregoing, this research was conceived and geared toward investigating some physical and chemical characteristics of Vertisols in Sudan Savannah Zone (SSZ) of Nigeria in order to furnish site specific data and management practice as well as suitable crop choices relevant to the farmers in the research locality for optimum exploitation of its potentials.

### MATERIAL AND METHOD

#### Description of Area of the study

Jigawa State (Longitude 12. 446°N, Latitude 9.7233° E), (Fig. I) has great potentials for agricultural production required for household consumption, commerce and industrial uses. About 70 % of the state's landmass is considered arable which makes it one of the most agriculturally endowed state in Nigeria (Kaugama and Ahmad, 2014). The eastern part of the state (Longitude 12. 14 – 12. 69° N, Latitude 9. 54 – 10. 21° E) is regarded food basket of the state (Wakili, 2016) (Fig. II). The largest area of Vertisols coverage is cropped to rice rotated to wheat and or green vegetables.

#### Geology, vegetation, hydrology and climatology of the area

The geologic formation of the upstream area consists of largely impermeable Basement Complex Rocks while the lowland area is of alluvial sediments of Chad formation. Rainfall period is from June to October, averaged 600 – 700 mm. Evapotranspiration ranges 1200 – 1500 mm (Sobawole, et al., 2001; Sanyu, 1994) and annual mean temperatures record 12 – 40 °C (Roger, 2013). The relief is characteristic of gentle slope. Basic soils of the wetlands are halomorphic and formed of deltaic alluvium. Whereas, the uplands surrounding the wetlands are non – leached ferruginous dune soils (Adams, et al., 1993). Composition of the vegetation entails diverse trees and grass species: *Guiera* spp and *Mitragyna* are found on low hills, while *Vetiveria* and *Andropogon gayanus* dominate swamp grassland. There are now large expanses of *Azadirachta indica*, *Adonsonia digitata* in many areas and large clusters of coppiced *Hyphaene thebaica* (doum palm) dominate great tracts of land. Many of the swamp grasslands have been succeeded and channels are blocked by an invasive weed species, *Typha dominguensis*.

#### Sampling frame

The sampling covered the entire eastern part where the Vertisols is dominant ((Long. 12. 14 – 12. 69° N, Lat. 9.54 – 10. 21° E). First, the area was divided into four sampling locations; the east, west, north and south. Ten (10) replicates each of composite and core samples were randomly collected (adopting diamond

shaped pattern) from each location to allow for complete coverage. A grand total of 120; (10 x 3 x 4) each of composite and core sample were thus collected for three depths considered; 0 – 30 cm, 30 – 60 cm and 60 – 90 cm. Samples were prepared and stored under standard procedure and subjected to appropriate laboratory analyses.

#### Laboratory Analysis

Auger (disturbed) samples, were air dried and passed through 2 mm sieve, for determination of particle size distributions by Bouyoucos hydrometer (Gee and Bauder, 1986); soil acidity (pH) in water and CaCl<sub>2</sub> as described by (Thomas, 1996); organic carbon by the dichromate wet oxidation method, (Nelson and Sommers, 1982); total nitrogen by the Kjeldahl digestion method (Bremner, 1982); available P by Bray No. 1 method (Bray and Kurtz, 1945), exchangeable bases (Black, (ed) 1965), exchangeable acidity by titration method (McLean, 1965), cation exchange capacity neutral K – acetate saturation and neutral NH<sub>4</sub><sup>+</sup> acetate Displacement method (Black, 1975) as described by the International Institute of Tropical Agriculture IITA 1979. ). Sample that passed through 5mm sieve was used to determine wet and dry mean weight diameters (Van Bavel, 1950) as modified by Kemper and Rosenau, (1986), Collected Undisturbed core (height = 5 cm, diameter = 4 cm and volume = 62.86 cm<sup>3</sup>) samples were used to determine dry bulk density (BD) (Anderson and Ingram, 1993), saturated hydraulic conductivity (Ksat) by constant head permeameter. Macro, micro and total porosities were determined as described by (Jarvis, et al., 2002; Dexter and Czyz, 2007; Reynolds, et al., 2008; Dexter, et al. 2008) as follows:

Pt = the volume of water at saturation (0.0 bar), Thus

$$Pt = \frac{W_s}{W_d} \times \frac{BD}{\rho_w} \quad \text{equation (1)}$$

Where:

W<sub>s</sub> = weight of saturated soil

W<sub>d</sub> = weight of oven dry soil

BD = Bulk density

ρ<sub>w</sub> = density of water at 20°C

Macro – porosity was computed using the formula:

$$P_{mac} = \theta_s (\psi=0.0 \text{ bar}) - \theta_m (\psi=0.33 \text{ bar}), \quad \text{equation (2)}$$

Where;

θ<sub>s</sub> = Moisture content at saturation,

θ<sub>m</sub> = Moisture content at field capacity,

ψ = Suction/ Pressure

Whereas, microporosity was determined as;

$$P_{mic} = P_t - P_{mac} \quad \text{equation (3)}$$

Where;

P<sub>t</sub> = Total porosity

P<sub>mac</sub> = Macro porosity

After obtaining aggregate size proportion of various mesh size the following formular was used to compute for mean weight diameters (MWD)

$$MWD = \sum_{i=1}^n x_i \omega_i \quad \text{equation (4)}$$

Where:

MWD = Mean Weight Diameter (mm),

n = number of aggregate fractions,

x<sub>i</sub> = arithmetic mean diameter of the size fraction and

ω<sub>i</sub> = proportion of the total water stable aggregates in the corresponding size fraction after deducting the weight of sand/gravels particles.

∑ = Summation.

(Van Bavel, 1950 as modified by Kemper and Rosenau, (1986). Sand correction was done.

Proportional macro aggregate and micro aggregate fractions were computed as described by (Ontel et al., 2015):

$$\frac{\text{Weight of stable macro aggregate (> 0.25 mm)}}{\text{Total weight of all the aggregate (200g)}} \quad \text{and}$$

$$\frac{\text{Weight of micro aggregate fraction (<0.25)}}{\text{Total weight of all the aggregates (200g)}}$$

respectively

equation (5) and (6)

#### Statistical analysis

GenStat (ed 4<sup>th</sup>) was used to run analysis of variance (ANOVA) in order to compare means for the three depths. While, means with significant difference were separate with the least mean square difference (LSD).

## RESULT AND DISCUSSION

### Chemical Properties

Based on table 1, the soil pH (H<sub>2</sub>O) was generally slightly acidic to moderately alkaline (6.11 – 7.90). It was higher at 95 % confidence level in 60 – 90 cm. Similar trend was also recorded with respect to pH in CaCl<sub>2</sub> which ranged between 5.44 – 6.33. This could have occurred due to presence of more exchangeable sodium in the lower layer. This agreed with Virmani (2000); Leka et al. (2012) who reported pH of Vertisols ranges between 6.3 and 7.4 for the surface soils and 6.9 to 8.0 for the subsoils. Furthermore, the pH (H<sub>2</sub>O) was more variable than pH (CaCl<sub>2</sub>), this may be due the ability of CaCl<sub>2</sub> solution to mimic natural field condition.

The result showed that exchangeable calcium (Ca) depicted somewhat uniform distribution but increased slightly with increase in depth. However, no significant difference was observed among the three depths. It ranged 2.81 – 2.95 Cmol/kg and was more variable at 0 – 30 (CV = 3.91 %) Both exchangeable magnesium (Mg) and potassium (K) showed significant variation and were statistically higher at 0 – 30 cm, followed by 30 – 60cm and lastly 60 – 90 cm. The higher Mg and K observed in the surface may be ascribed to their dissolution from soil mineral due to waterlogging. Pal et al. (2012) obtained similar trend in India. The exchangeable sodium (Na) on the other hand assumed a reverse trend; meaning it was significantly higher at 60 – 90 cm compared to the two upper depths. This may be attributed to its successive

deposition overtime by fluctuating watertable through capillarity rise. The Na was also more variable throughout the depths compared to the others exchangeable bases tested. This may probably be due to fluctuation of soil moisture status and water table, which tend to simultaneously dissolve and deposit sodium ions respectively.

Based on Table 2, cation exchange capacity (CEC) increased uniformly with depth thus observed to be higher at the 60 – 90 cm and was rated low to medium. The linearity depicted by CEC with depth is probably related to increase in clay content. Pal *et al.* (2012) reported in consistent CEC pattern with depth. The electrical conductivity (EC), showered inconsistent result but significantly varied at 0.01% probability level and was higher at 0 – 30 cm, it went down at 30 – 60 cm but sharply increased at 60 - 90 cm. The observed sharp increase in EC in 60 – 90 cm may be in response to the Na content increase. However, there is no threat to salinity development across the depths. Elsewhere, Virmani, (2000) indicated that the soils are not saline, which conform well with this finding.

Three parameters viz; organic carbon (OC) (4.78 – 7.33 g/kg), organic matter (OM) (8.33 – 12.50 g/kg) and total nitrogen (TN) (0.56 – 1.11 g/ kg) presented in Table 1 behaved in similarity. This is because they are from common source. They were all higher in the surface soil, 0 – 30 cm and steadily decrease down the depth, hence, the 60 – 90 cm got least mean value accordingly. Customary/ traditional application of organic materials: animal dung, home wastes and plant leftover, may have resulted in the higher OC, TN and OM in the 0 – 30 cm soil. This is in harmony with finding by (Yule and Ritchie, 1980) who stated that A<sub>1</sub> horizon has higher OC, TN and generally decreased in the lower horizons.

#### Physical properties

Table 3 presented result for particle size distribution. Sand content showed statistical variation, 0 – 30 cm obtained highest value while 60 – 90 cm recorded least mean value. A sharp contrast of that was obtained with respect to clay content. However, the two lower depths (30 – 60, 60 – 90 cm) were statistically similar. Silt percentage was similarly higher in the deeper depths (24.06 and 32.44 g/kg for 30 – 60 and 60 – 90 cm respectively) and more variable at 60 – 90 cm. The higher sand content and lower clay recorded in the 0 – 30 cm may be attributed to gradual deposition of sand particles from

neighboring sites. Dudal, (1965) had similarly reported a gradual increase in clay content with depth. It was thought that the increase in clay content with depth is due not to argillation of clay (migration of clay) but not due to inheritance from parent material.

Result for macroporosity (P<sub>mac</sub>) consistently decreased with increase in depth. On the other hand, microporosity (P<sub>mic</sub>) showed nearly opposite outcome. Total porosity (P<sub>t</sub>) reflected zigzag/ broken trend; it decrease at 30 – 60 cm but plateaued at 60 – 90 cm. The two deeper depths showed increased bulk density (BD), whereas, the upper 0 – 30 cm recorded the least mean value statistically ( $P \leq 0.05$ ). The reduced BD may be due to pulverization through tillage operation and organic matter decomposition. This is in concord with Eswaran *et al.* (1988); Mermut *et al.* (1996) and Butler and Hubble, (1977). The bulk density of Vertisols usually tends to increase with depth, due to compression caused by overburden weight.

The 60 – 90 cm showed limited saturated hydraulic conductivity (K<sub>sat</sub>) comparative to the two overlying layers. This is attributed to dominance of micropores and surface sealing effect observed in the area. The dry mean weight diameter (MWD<sub>d</sub>) and macro aggregate proportion indices showed no significant difference across the three depths. Oppositely, the wet mean weight diameter (MWD<sub>w</sub>) significantly varied, in which 60 – 90 cm and 30 – 60 cm depicted similarity and were higher than the surface layer. The greater mean value in these lower depths may be due to less disturbance and more binding assist by clay. An irregular pattern was observed with respect to microaggregate proportion. The 0 – 30 cm and 60 – 90 cm were statistically not different while the 30 – 60 cm recorded lowest mean values.

#### CONCLUSION

An investigation was carried-out objectively to determine the productivity potential of Vertisols in Sudan Savannah Zone of Nigeria. Standard procedures for soil sample collection, processing and analyses were adopted. ANOVA was used for the analysis using GenStat ed 13<sup>th</sup>. From the result, it can be concluded that, the Vertisols is potentially productive and resilient to water and wind degradation. However, addition of both organic and inorganic fertilizer is necessary for attaining optimum agricultural productivity and profitability. Rice and wheat crop could perform well when grown on these soils.

**Table 1: Mean values for some chemical properties of the Vertisols at varying depth**

Depth	pH (H <sub>2</sub> O)	pH (CaCl <sub>2</sub> )	OC	OM	TN	Avail.P
				g/ kg		mg/kg
<b>0 – 30 cm</b>	6.11b	5.44c	7.33a	12.50 a	1.11a	2.39 a
<b>SD</b>	0.42	0.14	0.17	0.39	0.09	0.08
<b>CV (SD/X)</b>	6.87	2.74	2.32	3.12	8.11	3.35
<b>30 – 60 cm</b>	6.50b	5.78b	5.83b	10.17b	0.89b	2.11b

<b>SD</b>	0.25	0.17	0.21	0.18	0.07	0.09	
<b>CV(SD/X)</b>	3.85		2.94	3.60	1.77	7.87	4.27
<b>60 – 90 cm</b>	7.90a	6.33a	4.78c	8.33c	0.56c	1.68 c	
<b>SD</b>	0.131	0.79	0.132	0.288	0.085	0.091	
<b>CV(SD/X)</b>	1.66	12.48	2.76	3.46	15.17	5.42	

CV = coefficient of variance, SD = Standard deviation, X = mean, OM = organic matter, OC = organic carbon and TN = total nitrogen, Avail.P = available phosphorus. Means with the same letter across the column are not statistically different.

**Table 2: Mean values for some chemical properties of the Vertisols at varying depth**

Depth	Ca	Mg	K	Na	CEC	EC	EA
	cmol <sup>+</sup> /kg			dS/m			
<b>0 – 30 cm</b>	2.81	1.76 a	0.34a	0.39b	5.28c	0.061b	0.38
<b>SD</b>	0.11	0.04	0.01	0.03	0.13	0.004	0.05
<b>CV (SD/X)</b>	3.91	2.27	2.56	7.69	2.46	6.56	13.16
<b>30 – 60 cm</b>	2.96	1.53 b	0.27b	0.40b	6.05b	0.04c	0.48
<b>SD</b>	0.07	0.04	0.007	0.02	0.13	0.003	0.04
<b>CV (SD/X)</b>	2.36	2.61	2.59	5.00	2.15	7.5	8.33
<b>60 – 90 cm</b>	2.95ns	1.34 c	0.209c	0.613a	6.44a	0.138a	0.41ns
<b>SD</b>	0.064	0.045	0.008	0.043	0.120	0.015	0.065
<b>CV (SD/X)</b>	<b>2.17</b>	<b>3.36</b>	<b>3.83</b>	<b>7.01</b>	<b>1.86</b>	<b>10.86</b>	<b>15.85</b>

Ca = Calcium, Mg = Magnesium, K = Potassium, Na = Sodium, CEC = cation exchange capacity, Ec = electrical conductivity, EA = exchangeable acidity, SE = standard error, CV = coefficient of variability, SD = Standard deviation, X = mean. Means with the same letter across the column are not statistically different.

**Table 3: Mean values for some physical properties of the Vertisols at varying depth**

Depth	Sand	Silt	Clay	Textural class	Pmac	Pmic	Pt
	%						
<b>0 – 30 cm</b>	59.17a	19.72b	21.11b	SCL	29.44a	34.58b	64.02a
<b>SD</b>	2.39	2.45	1.03		1.86	1.66	2.15
<b>CV (SD/X)</b>	4.04	12.4	4.88		6.32	4.80	3.36
<b>30 – 60 cm</b>	51.00b	21.22ab	27.79a	SCL	23.61b	33.72b	57.33b
<b>SD</b>	1.62	2.04	0.95		1.97	1.91	1.52
<b>CV(SD/X)</b>	3.18	9.61	3.42		8.34	5.66	2.65
<b>60 – 90 cm</b>	43.5 c	24.06a	32.44a	SCL	19.11c	42.19 a	61.30ab
<b>SD</b>	2.132	2.151	1.5215		0.632	1.998	2.761
<b>CV (SD/ X)</b>	4.9	8.94	4.69		3.33	4.74	4.5

Pmac = macro porosity, Pmic = micro porosity, Pt = total porosity, SE = standard error, CV = coefficient of variability, SD = Standard deviation, X = mean. Means with the same letter across the column are not statistically different.

**Table 4: Mean values for some physical properties of the Vertisols at varying depth**

Depth	BD	Ksat	MWDd	MWDw	Mac.Agg	Mic. Agg.	
	mg/ m <sup>3</sup>	mm/s			%		
<b>0 – 30 cm</b>	1.00b	2.89a	1.01	0.38b	19.67	78.91a	
<b>SD</b>	0.01	0.14	0.04	0.03	1.60	3.63	
<b>CV(SD/X)</b>	1	4.84	3.96	7.89	8.13	4.60	

<b>30 – 60 cm</b>	1.20a	2.56a	1.10	0.42 a	22.67	75.67b
<b>SD</b>	0.05	0.09	0.03	0.04	2.12	5.533
<b>CV (SD/X)</b>	4.17	3.52	2.73	9.56	9.31	7.31
<b>60 – 90 cm</b>	1.25a	1.83b	1.13ns	0.40a	19.50 ns	79.00a
<b>SD</b>	0.045	0.072	0.033	0.0071	1.181	1.222
<b>CV (SD/X)</b>	3.60	3.93	2.92	1.77	6.06	1.55

BD = bulk density, Ksat = saturated hydraulic conductivity, MWDD = dry mean weight diameter, MWDw = wet mean weight diameter, Mac.Agg = macro aggregate, Mic. Agg. micro aggregate, CV = Coefficient of variance, SD = Standard deviation, X = mean. Means with the same letter across the column are not statistically different.

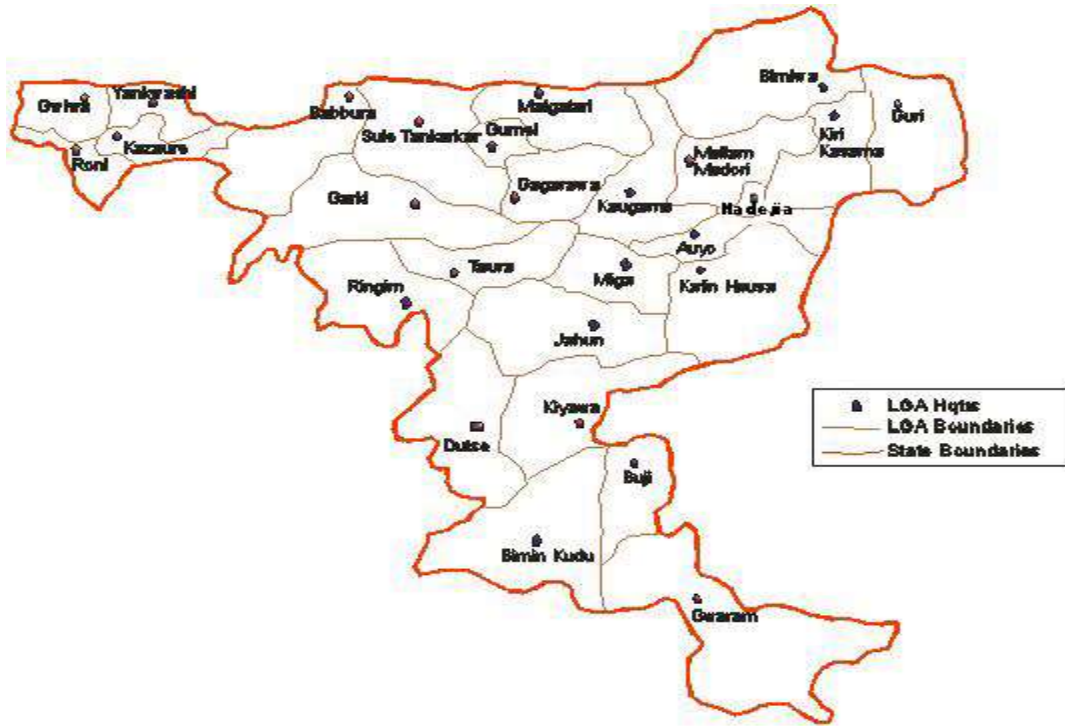


Figure I: Map of Jigawa State, Nigeria

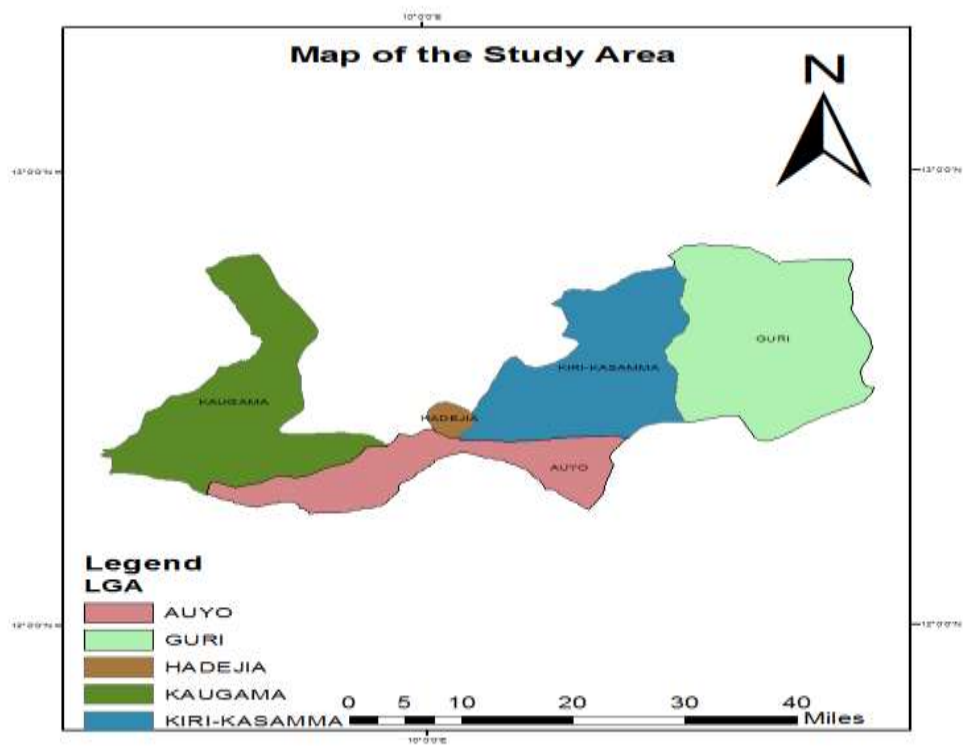


Fig. II: Location of the Study Area

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